

GEOPHYSICAL INVESTIGATION  
OF A DIABASE DIKE

A THESIS

Presented to

The Faculty of the Division of Graduate  
Studies and Research

by

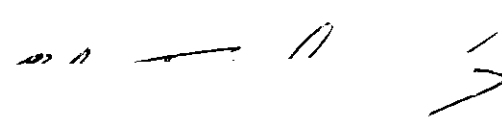
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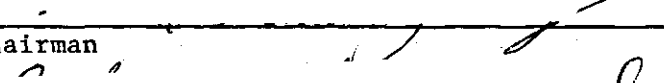
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Master of Science in Geophysical Sciences



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GEOPHYSICAL INVESTIGATION OF A DIABASE DIKE

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## SUMMARY

Analysis of gravity, aeromagnetic, ground-level magnetic, paleomagnetic, and high altitude infrared data has revealed a complex injection zone of diabase in central Meriwether County, Georgia. This zone, which previously was considered to be a single large (30-40 meters wide) diabase dike, has been found in places to reach a width of one kilometer.

A simple Bouguer gravity map of the area shows the dikes, which compose the injection zone, to be responsible for an anomaly in the regional gravity trend of about one to two milligals. Combined results from five gravity and three total-field magnetic traverses has suggested the dikes to dip at about  $70^{\circ}$  toward  $N70^{\circ}E$ . The mean of the observed ground-level magnetic anomalies for the dikes is a 1000 $\gamma$  positive anomaly (total field).

On the basis of observed paleomagnetic data and computed bulk susceptibilities the ratio of the remanent to induced magnetizations (Koenigsberger's ratio,  $Q$ ) has been calculated to be between 0.25 and 0.5. However, the shapes of the observed ground-level magnetic anomalies suggests a  $Q$  value of 1.0 or slightly higher for the Meriwether dike.

The total-field magnetic map based on already existing aeromagnetic data has been recontoured under the assumption that the Meriwether dike is continuous, as suggested by ground-level data taken during this investigation.

On the recontoured map the dike is the most prominent magnetic

feature of the area which suggests that the spacing and direction of the aeromagnetic flight lines is inappropriate for correlating dike caused anomalies of limited extent, even though the dike anomalies are large in magnitude.

High altitude infrared photographs show a change in intensity of reflected infrared radiation from vegetation growing on diabase derived soils suggesting that such pictures can be used to map the outcrop pattern of the dike in heavily wooded, highly weathered, and other inaccessible areas. It is suggested that this technique be applied to other areas of eastern North America to determine outcrop patterns of the dikes in areas not yet geologically mapped.

## CHAPTER I

### INTRODUCTION

A system of diabase dike swarms outcrops throughout eastern North America from Alabama to Nova Scotia (Cohee, et al., 1962), (Stockwell, et al., 1969). Southward along the Appalachian belt from Nova Scotia through New England and into the southern Piedmont, a gradual systematic change in strike is evident (Figure 1). King (1961) commented that the pattern of diabase dikes in eastern North America is probably the result of deep-seated stresses, but that the cause of the stresses is not apparent. May (1971) noted that the systematic pattern of the dikes in eastern North America is actually part of a larger, radial, pattern of diabase dikes surrounding the North Atlantic (Figure 1) and suggested that this pattern is a result of the stress field associated with the onset of North Atlantic sea-floor spreading in Late Triassic or Early Jurassic time.

The age of the dikes in eastern North America is currently in dispute. On the basis of fossil pole positions, de Boer (1967) suggested a Jurassic age for the dikes. On the other hand, Armstrong and Besancon (1970) found whole rock K-Ar ages to fall into two clusters: one at about 200 million years (Middle Triassic) and another smaller one between 225 and 230 million years (Early Triassic). Armstrong and Besancon (1970) suggested that the older dates may represent the actual time of emplacement, while the younger group may be indicative of "burial"

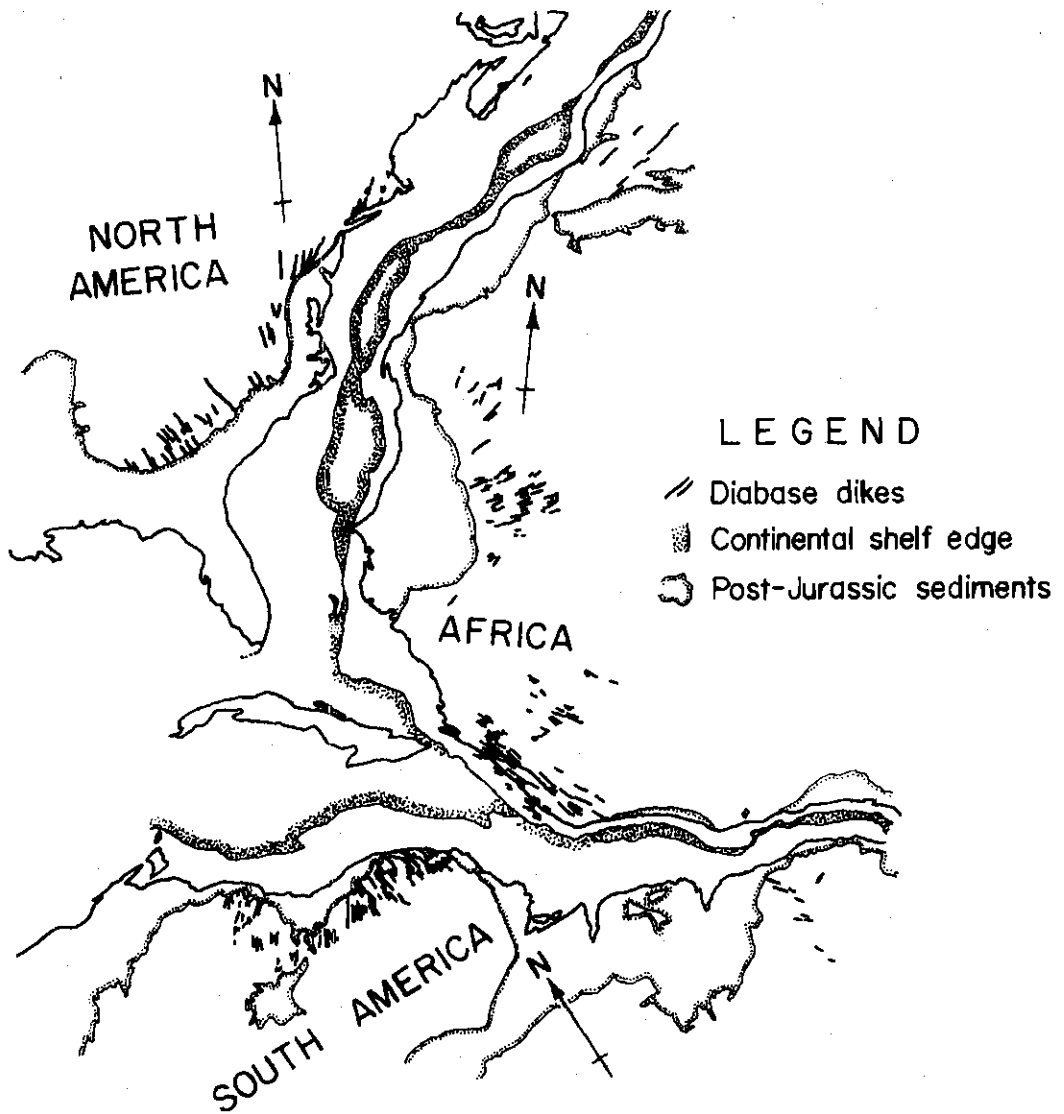


Figure 1. Map of Diabase Dikes in Eastern North America, West Africa, and Northeastern South America with the Continents Restored to Their Relative Position in the Triassic (after May, 1971). (North arrows indicate present geographic direction for each continent.)

metamorphism.

The chemistry of the diabase dikes, previously thought to be uniform (Lester and Allen, 1950) has been shown by Weigand and Ragland (1970) to vary in both major and trace element assemblages. Weigand and Ragland (1970) divided the dikes of eastern North America into several provinces based on the occurrence of three main chemical types: (1) olivine-normative, (2) high-TiO<sub>2</sub>, quartz-normative and (3) low-TiO<sub>2</sub>, quartz-normative.

The tracing of individual dikes along strike and the determination of thickness and dip by the usual geological field methods is inhibited in most areas of Southeastern North America by intense chemical weathering and dense vegetation. However, Johnson and Watkins (1963) noted a coincidence of dike outcrops and low-amplitude (less than 20% at an altitude of 800 feet above ground level), positive magnetic anomalies along aeromagnetic flight lines in north-central Virginia. On the basis of linear trends in the occurrence of some of these magnetic highs, Johnson and Watkins (1963) inferred the extension of dikes into areas where no outcrops are known. The success of Johnson and Watkins (1963) suggests that geophysical methods at closer range, i.e. aeromagnetics at 500 feet, ground-level magnetics and detailed gravity, could yield detailed structural information as well as outcrop patterns. The purpose of this study was to apply these geophysical methods to a study of the structure of a diabase dike.

A portion of the long dike in central Meriwether County in west-central Georgia (hereinafter referred to as the Meriwether dike) was selected for study. The dike extends southward from an area north-east

of Newnan in Coweta County, to an area south of the Towaliga fault in southern Meriwether county (Figure 2). This area of study was selected because of minimal secondary geophysical disturbances, such as gravity and magnetic anomalies associated with fault zones and basic intrusives other than the diabase dike. In addition, the roads in this area traverse the Meriwether dike at almost right angles and are spaced at approximately two miles, affording good access for a geophysical study.

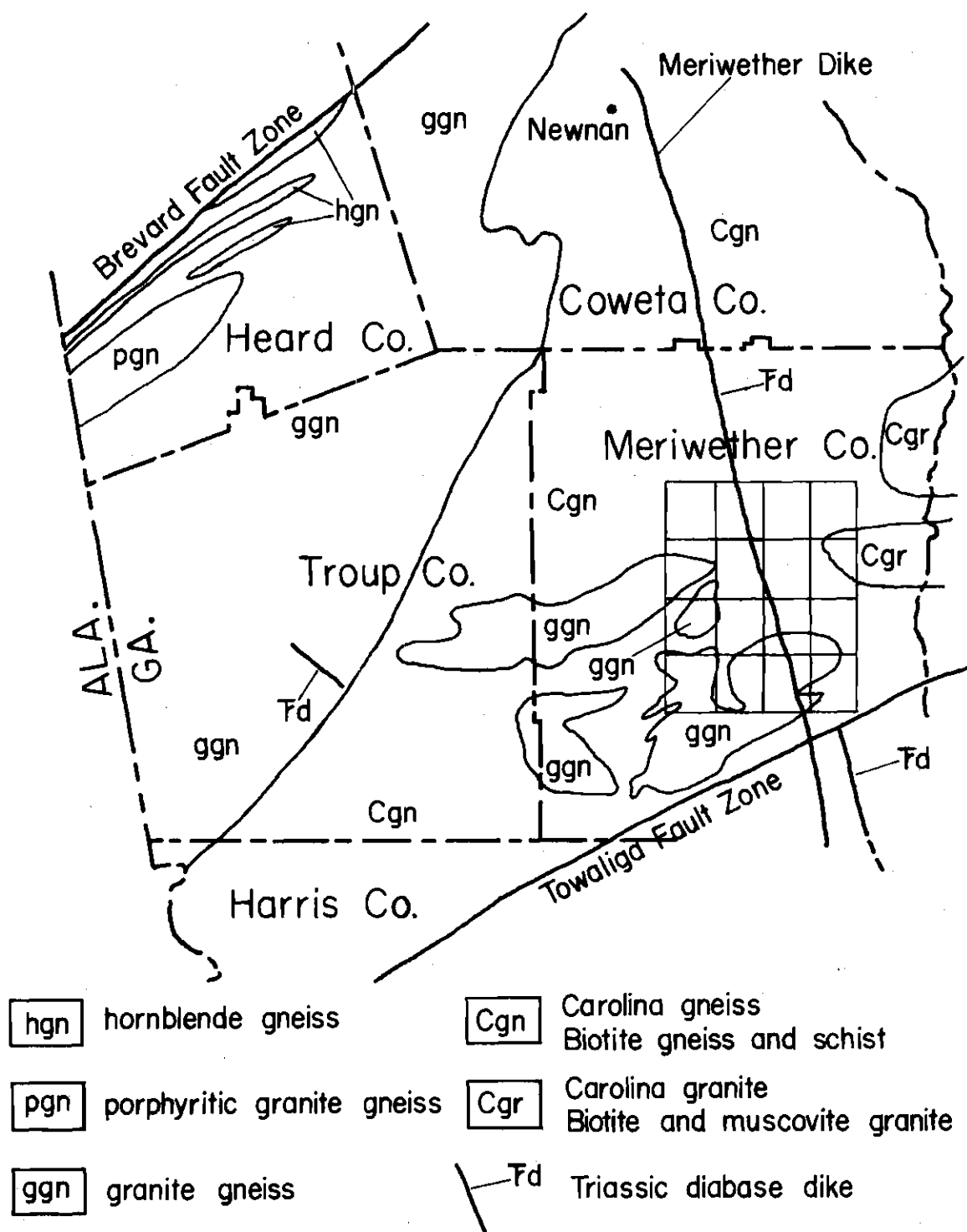


Figure 2. Geology of the Inner Piedmont of West Georgia (after the 1939 Geologic Map of Georgia). (Hatched square grid shows area of detail study.)

## CHAPTER II

### GEOLOGICAL SETTING AND REGIONAL GEOPHYSICS

#### Geologic Setting

According to the Geologic Map of the State of Georgia (Georgia Department of Mines, Mining and Geology, 1939) the geologic formations of the study area consist of granites intermingled with both biotite and granite gneisses (Figure 2). Recent reconnaissance mapping by R. D. Bentley (personal communications) has revealed the geology of the Inner Piedmont (between the Brevard and Towaliga Fault Zones) to be more complicated than suggested by the 1939 Georgia Geology Map. Bentley and Neathery (1970) renamed the rocks of the southeastern portion of the Inner Piedmont the Opelika Complex and divided it into two stratigraphic units, the Loachapoka Schist and the Auburn Gneiss-schist (Figure 3). These units were interpreted by Bentley (Bentley and Neathery, 1970) as metamorphosed sediments which are extensively intermingled with granites. The Meriwether dike strikes about  $N20^{\circ}W$  through the area and intersects the rocks of the Opelika Complex at about  $60^{\circ}$ .

#### Regional Gravity

Based on measurements taken at 195 stations, a simple Bouguer gravity map of the area under investigation was constructed (Figure 4). The procedures used for the reduction of the data and estimating errors in the final values are given in Appendix I. Estimated error in the regional data is  $\pm 0.35$  milligals. Isogals trend approximately north-



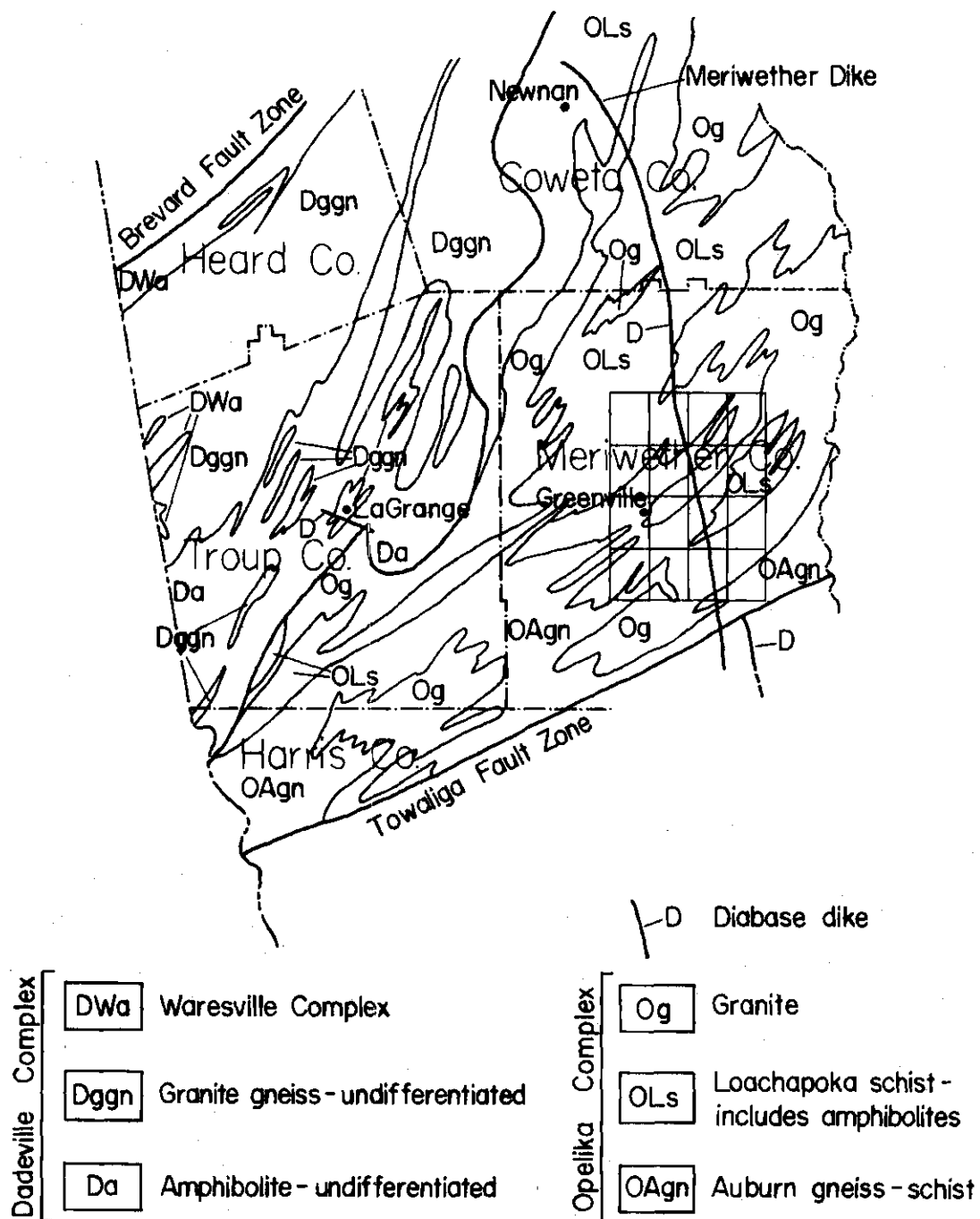


Figure 3. Geology of the Inner Piedmont of West Georgia (after Bentley Neathery, 1970). (Hatched square grid shows area of detail study.)

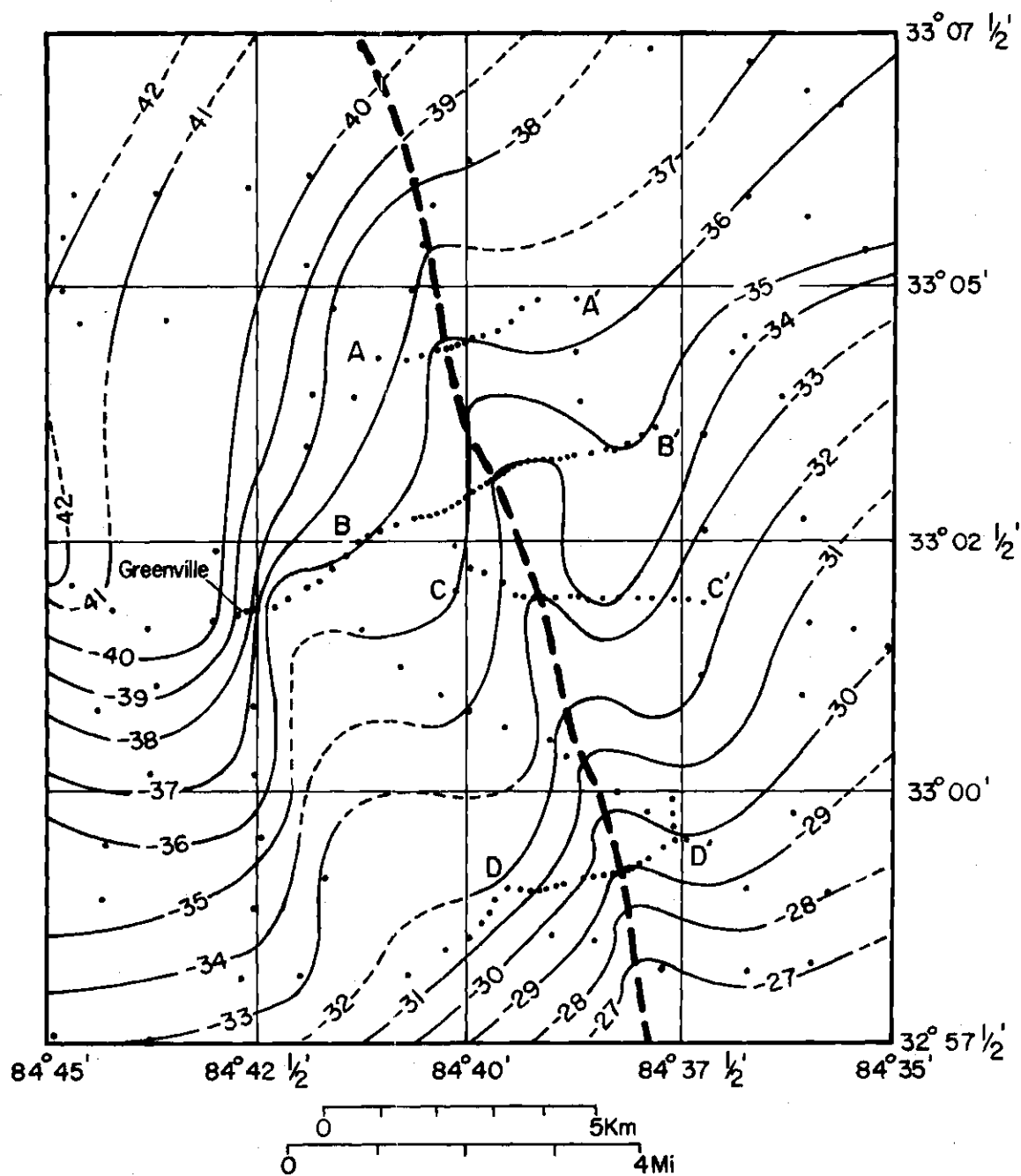


Figure 4. Simple Bouguer Gravity Map of Detail Study Area. (Dashed line gives location of Meriwether dike. Letters designate beginning and end of detailed gravity lines.)

east-southwest, and the regional gradient is about -0.8 milligals/kilometer to the northwest. The prominent features on the simple Bouguer gravity map are the steep gradient in the west-central portion of the study area and an anomaly in the regional trend of one to two milligals beginning in the southeastern part of the study area and continuing northward through the area. The steep gradient is coincidental with the contact between the Loachapoka Schist and granite west of Greenville (Figure 3). The anomaly in the regional trend is largely caused by the existence of the Meriwether dike. Too few gravity stations were taken in the extreme northern portion of the study area to warrant contouring, except by extrapolation. The anomaly probably extends northward out of the study area coincident with the dike.

#### Regional Magnetism

The North-Central Georgia Aeromagnetism Map (U.S. Geological Survey, open file, 1973) partially fills the gap (Figure 5) between the two previous aeromagnetic surveys of North Georgia Nuclear Laboratory (Philbin, Petrafeso and Long, 1964) and Savannah River Plant (Petty, Petrafeso and Moore, 1965). The area of this investigation (Figure 6) is part of the area covered by this latest work.

Similar to the regional gravity, magnetic features trend NE-SW. Lack of steep gradient immediately east of Greenville suggests that the units responsible for the observed steep gravity gradient in this area have a low contrast in magnetic susceptibility. The most conspicuous feature is the string of localized magnetic highs trending from north to south and having centers located at intersections of the dike and the

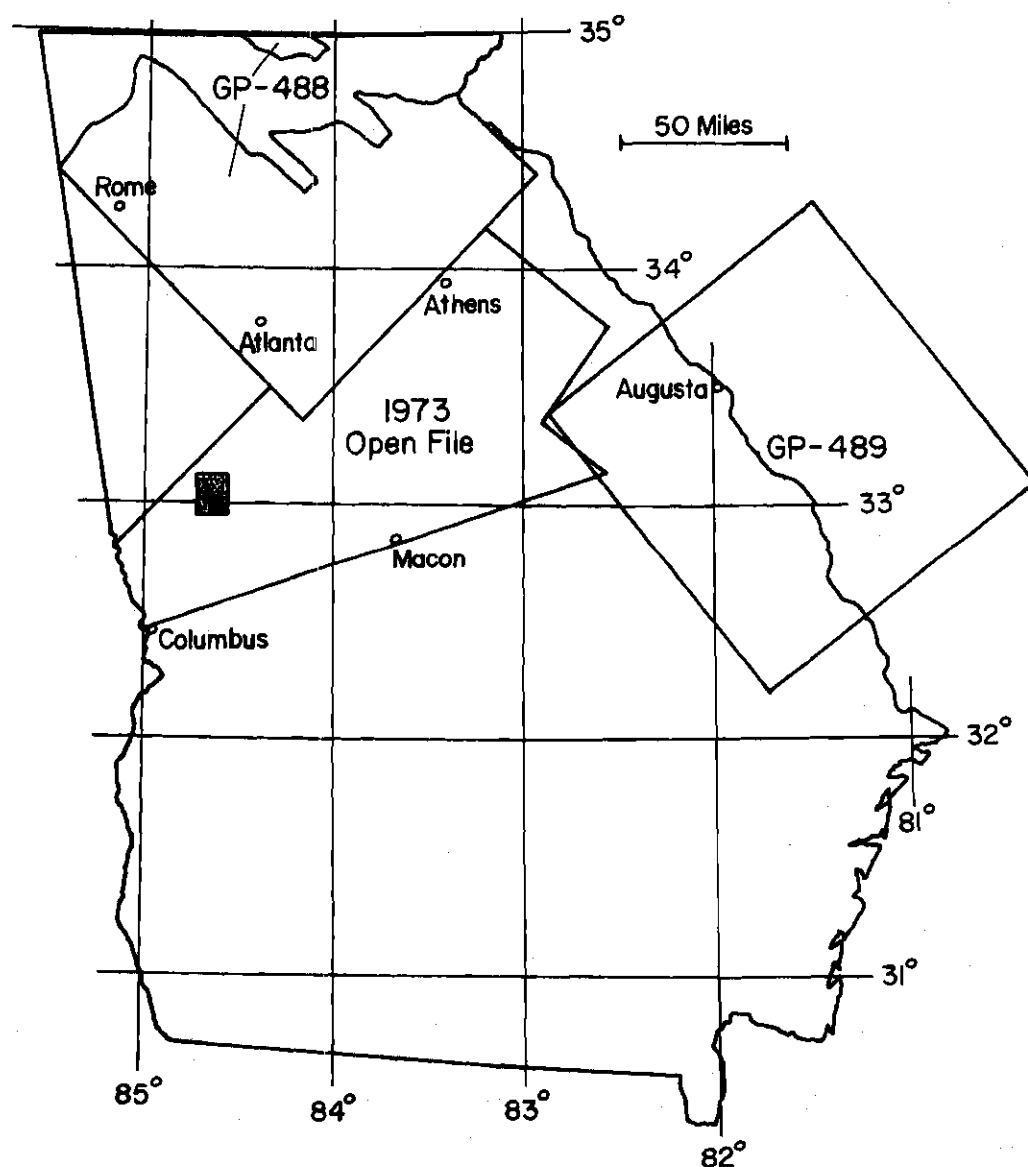


Figure 5. Index Map of Georgia Showing Location of Detail Study Area and Existing Aeromagnetic Maps. (Maps GP-488, North Georgia Nuclear Laboratory, and GP-489, Savannah River Plant, have been published by the U. S. Geological Survey. The North-Central Georgia Map is available on open file at the Georgia Geological Survey. Stipled rectangle is area of detail study.)

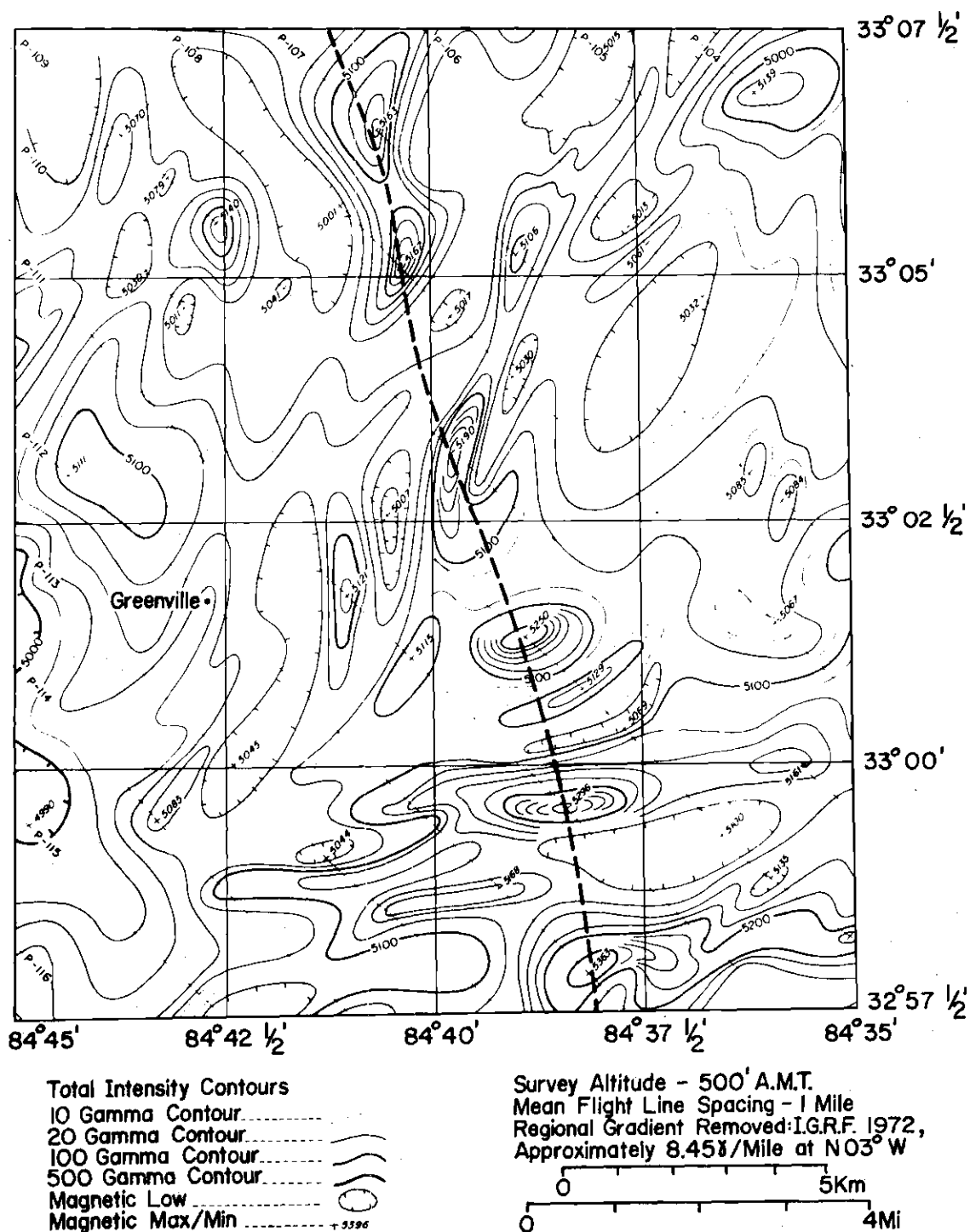


Figure 6. Portion of North-Central Georgia Aeromagnetic Map Covering Area of Detail Study. (U.S.G.S. Open File, 1973) (Dashed line gives location of Meriwether dike. Anomalies are with reference to 50,000γ datum.)

flight lines. These magnetic highs are probably a result of the presence of the Meriwether dike.

The Meriwether dike, which outcrops on roads along its entire length, is probably more continuous than suggested by the series of magnetic highs in Figure 6. Detailed ground-level magnetic traverses of the Meriwether dike at points between aeromagnetic flight lines (Figure 7) support this supposition of continuity. Prominent (1000 $\gamma$ ) magnetic highs coinciding with observed dike outcroppings (Figure 8) suggested a recontouring of the aeromagnetic data.

Since the continuous flight line data were not available, all points where flight lines intersect contour lines were assumed to represent true values (the data were extrapolated by computer methods between flight lines). Latitude, longitude, and the total magnetic intensity were determined for each intersection and these points were plotted for recontouring. Under the assumption that the Meriwether dike is continuous along strike, straight lines were drawn connecting points of equal total magnetic intensity on adjacent flight lines in the region where the dike outcrops. In other areas contour lines were assumed to be influenced by the regional geology (after Bentley and Neathery, 1970).

Although no magnetic interpretation is unique, this one (Figure 9) does strongly suggest that the magnetic anomalies associated with the dike can be represented by a continuous linear anomaly. This interpretation is necessary for two-dimensional analysis of the dike by comparison of its observed anomalies with those of easily computed models.



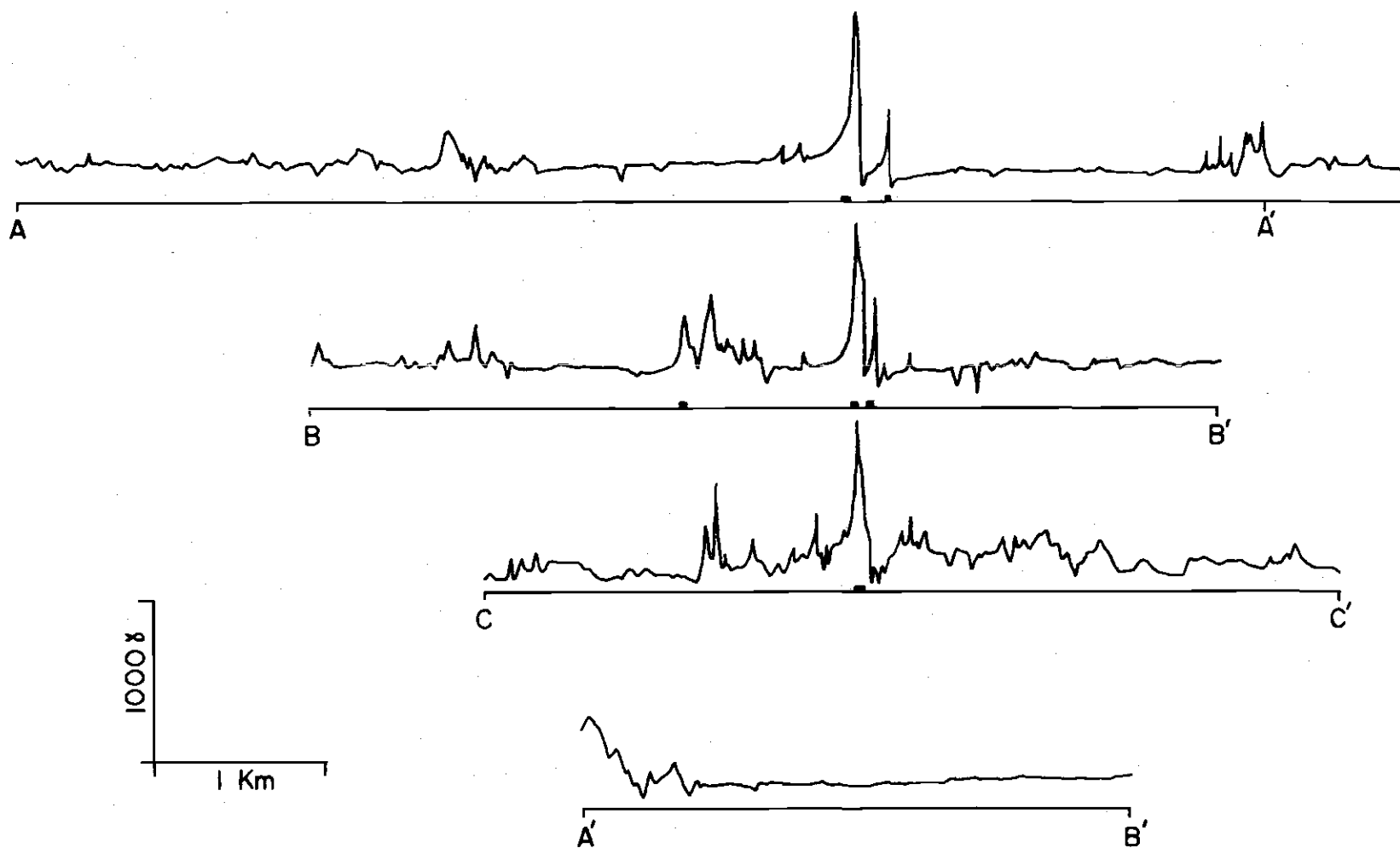


Figure 8. Ground-Level Magnetic Profiles (Total Field). (Heavy bar on profile base indicates diabase outcrop. See Figure 7 for location of profiles.)



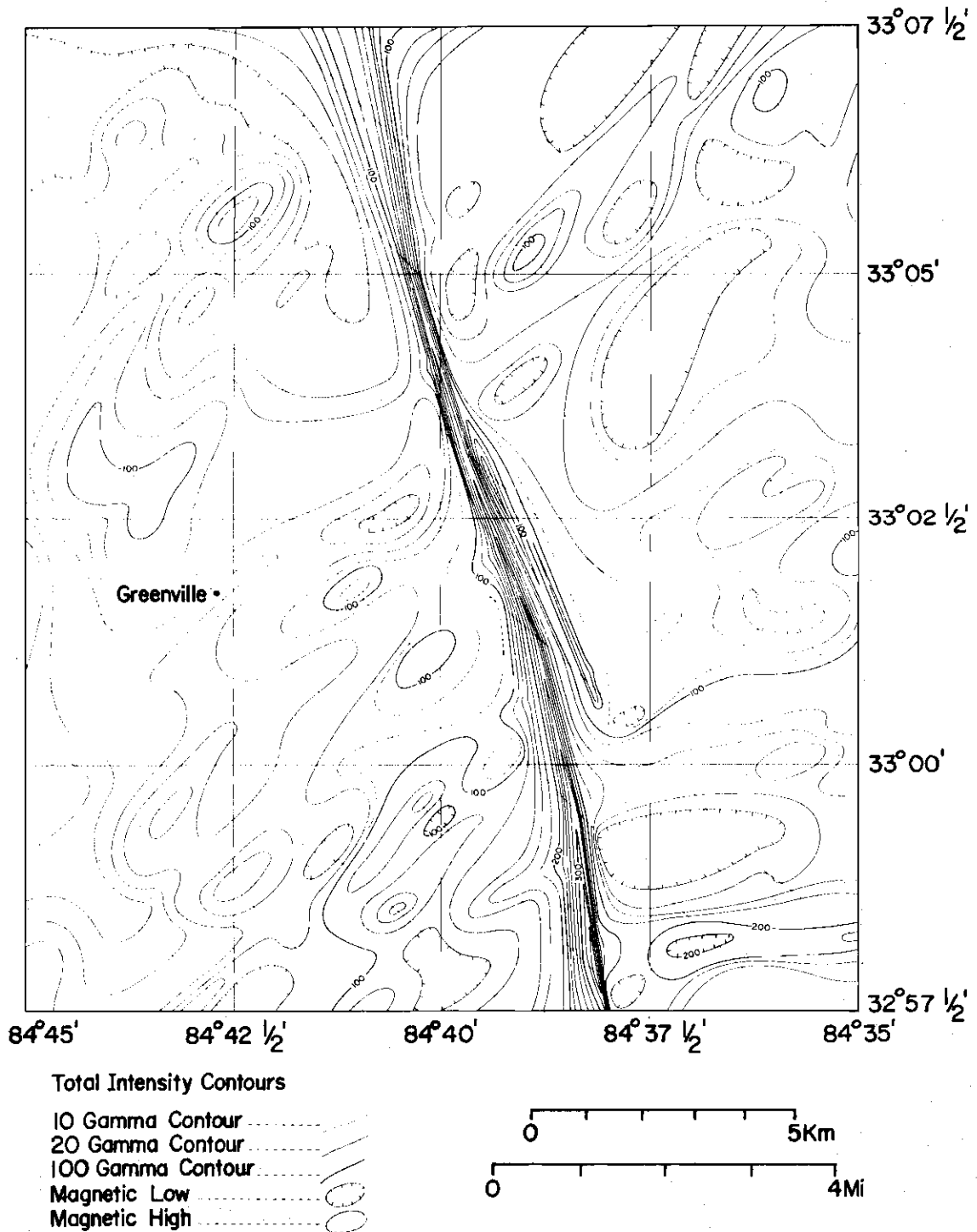


Figure 9. Aeromagnetic Map (Total Field) of the Study Area Recontoured Assuming the Meriwether Dike to be a Continuous Magnetic Feature. (Anomalies are with reference to 55,000 $\gamma$  datum.)

### CHAPTER III

#### PHYSICAL PROPERTIES OF THE MERIWETHER DIKE

##### Density

Diabase lends itself especially well to geophysical investigation. The density of the Meriwether diabase is  $3.0 \pm 0.05 \text{ gm/cm}^3$  (as determined by gravimetric methods applied to three field samples). Watson (1902) determined the densities of the Loachapoka Schist and local granites to be  $2.64 \pm 0.01 \text{ gm/cm}^3$  and  $2.70 \pm 0.01 \text{ gm/cm}^3$  respectively. Thus, the density contrast between the diabase of the Meriwether dike and the surrounding rocks (average density  $2.67 \pm 0.02 \text{ gm/cm}^3$ ) is  $0.33 \pm 0.07 \text{ gm/cm}^3$ . A vertical dike 30 meters wide (average observed outcrop width of the Meriwether dike) with this density contrast should yield a vertical gravity anomaly of approximately 1.5 milligals, an anomaly easily discerned with contemporary instrumentation (see Appendix I).

##### Magnetic Susceptibility

Several attempts have been made to statistically relate the bulk susceptibility of rocks to petrological parameters. One such attempt was that of Balsey and Buddington (1958) who related the susceptibility of some Adirondack rocks to the fractional volume of all the minerals visually identified as magnetite. This volume would generally include any Fe-Ti oxide minerals of spinel structure. Their empirical formula for the bulk susceptibility ( $k$ ) in C.G.S. units is

$$k = 2.6 \times 10^{-3} V^{1.11} \quad (1)$$

where V is the volume percentage of all minerals visually identified as magnetite.

Petrographic examination of the Meriwether dike (Lee, 1971) showed that opaque grains (assumed to be Fe-Ti oxides of spinel structure) comprise 2.3 to 4.3 percent of the rock with a higher percentage occurring in the finer-grained zones near its edges. Using the empirical formula of Balsey and Buddington (1958) (Equation 1) and the range of volume percentages of opaques given by Lee (1971), the bulk susceptibility of the diabase composing the Meriwether dike was estimated to be between 0.0066 and 0.0130 cgs.

## CHAPTER IV

### DETERMINATION OF DIP ANGLE FROM GRAVITY AND MAGNETIC ANOMALIES

#### Theoretical Models

Theoretical anomalies were calculated for sets of models to examine the effects of dip angle on expected gravity and magnetic anomalies. The method of Talwani, Worzel, and Landisman (1959) was used to compute the gravity anomalies that would be expected for dikes dipping at various angles. Because the observed outcrop width would be fixed for a particular dike, anomalies were computed for models with an outcrop width of 30 meters dipping at 90, 75, 60, and 45° (Figure 10). As would be expected, the gravity anomaly caused by a vertical dike is symmetric about the center of the dike. However, as the dip angle decreases the anomaly becomes asymmetric with the peak shifting toward the direction in which the dike is dipping. The anomaly tails off more slowly on the side toward which the dike dips than the other. Even though there are obvious differences in the anomalies for various values of dip angles, a superimposed regional trend may make determination of the dip more difficult than suggested by Figure 10; in fact, unless the regional trend is simple and well defined, the determination of the dip may be impossible.

Although the dip angle has an obvious effect on the shape and asymmetry of the total-field anomaly (Figure 11), the effects of natural

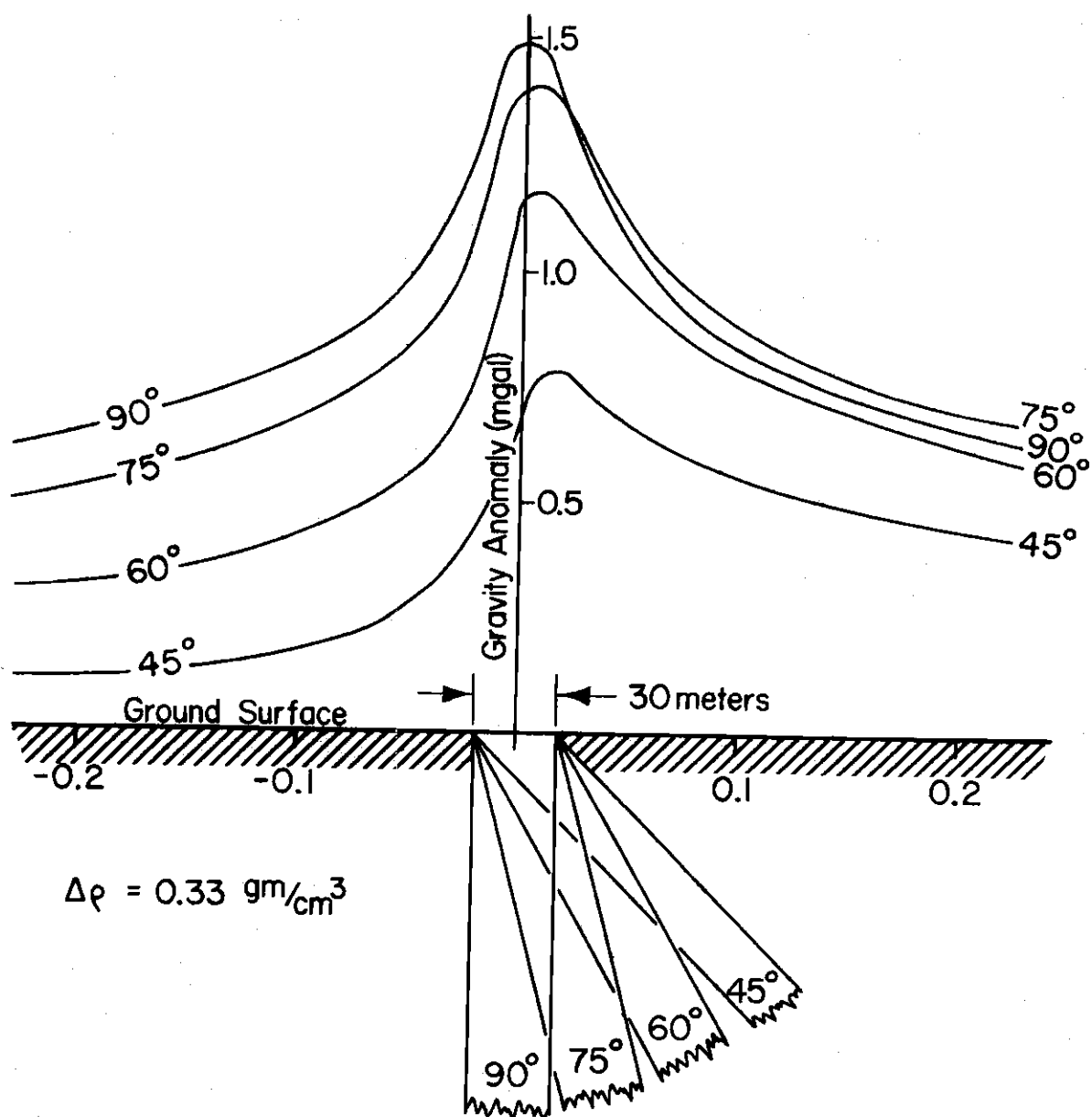


Figure 10. Calculated Curves Showing Effect of Dip Angle on Vertical Gravity Anomaly.

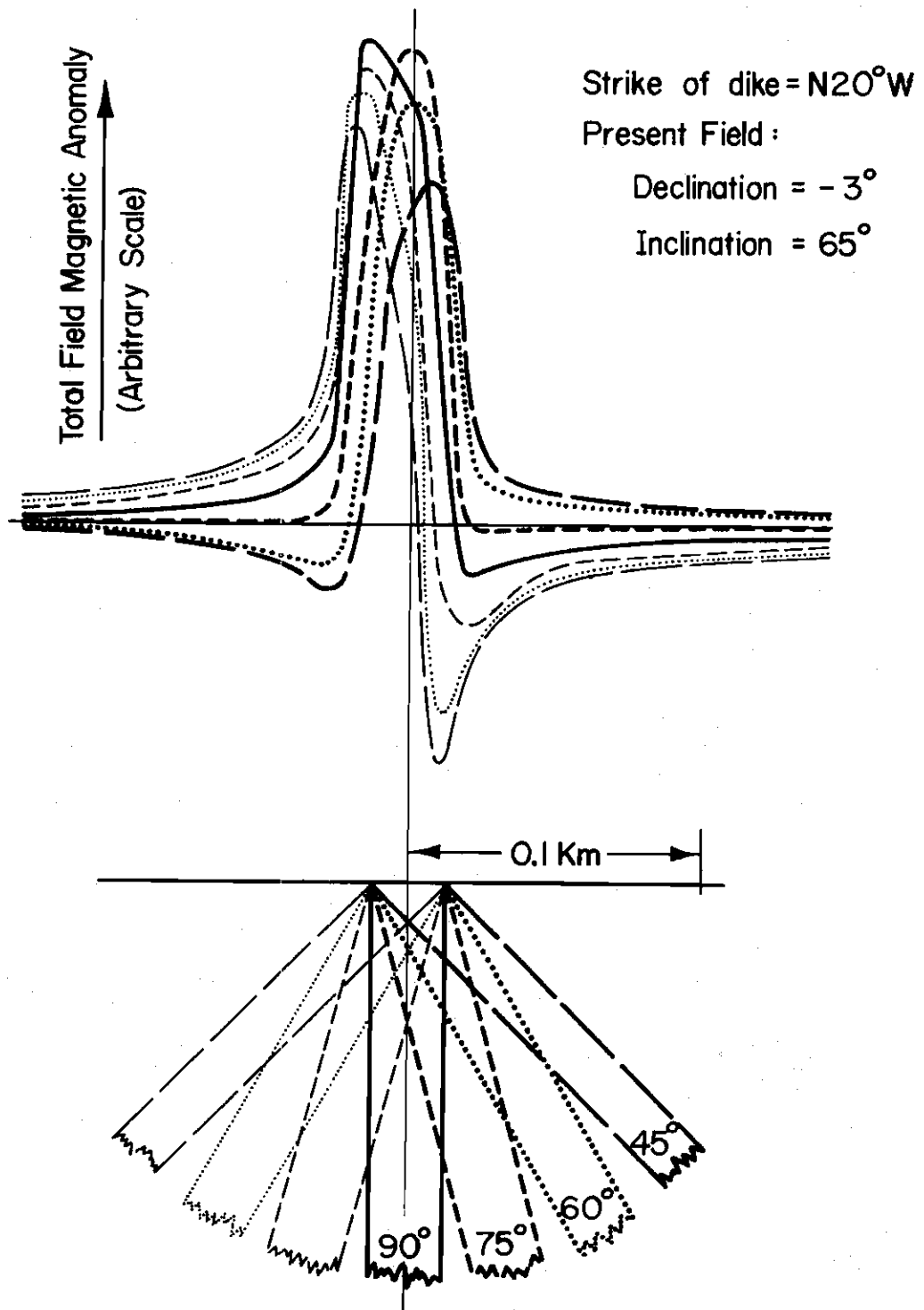


Figure 11. Calculated Curves Showing Effect of Dip Angle on Total Field Anomaly Due to Induction Magnetization Only.

remanent magnetization, when added to the induced magnetization, may complicate the determination of dip angle from the observed anomaly (Hood, 1963). For this reason, the angle of dip was determined using observed gravity anomalies only. The magnetic anomalies, because of their sharpness as compared to the gravity anomalies (Figures 8 and 10), were used only for lateral location of the dike where outcrops were nonexistent.

#### Detailed Gravity Profiles

Five detailed gravity lines with an average station spacing of 0.15 kilometers and a total length of 20 kilometers were obtained. Station spacing was decreased to about 0.075 kilometers in the immediate proximity of the dike. Four lines traverse the Meriwether dike in the prescribed study area (Figure 4) and the fifth traverses it along Georgia State Highway 16 southeast of Newnan and north of the study area. The observed anomalies (Figures 12 through 16), when smoothed, are about 1.0 kilometer wide at half their maximum value. This width is considerably greater than the computed anomaly width for a vertical dike 30 meters wide, approximately 0.1 kilometers at half its maximum value (Figure 10). Numerous individual peaks superimposed on the broad peaks suggest that a 0.75 kilometer wide swarm of dikes, rather than a single dike, is responsible for the observed anomaly. Such a hypothesis is further supported by the ground-level magnetic profiles, (Figure 8).

#### Profile A-A'.

Gravity line A-A' (Figure 4) was established along a dirt road about 3.0 kilometers north of Georgia State Highway 109. The observed

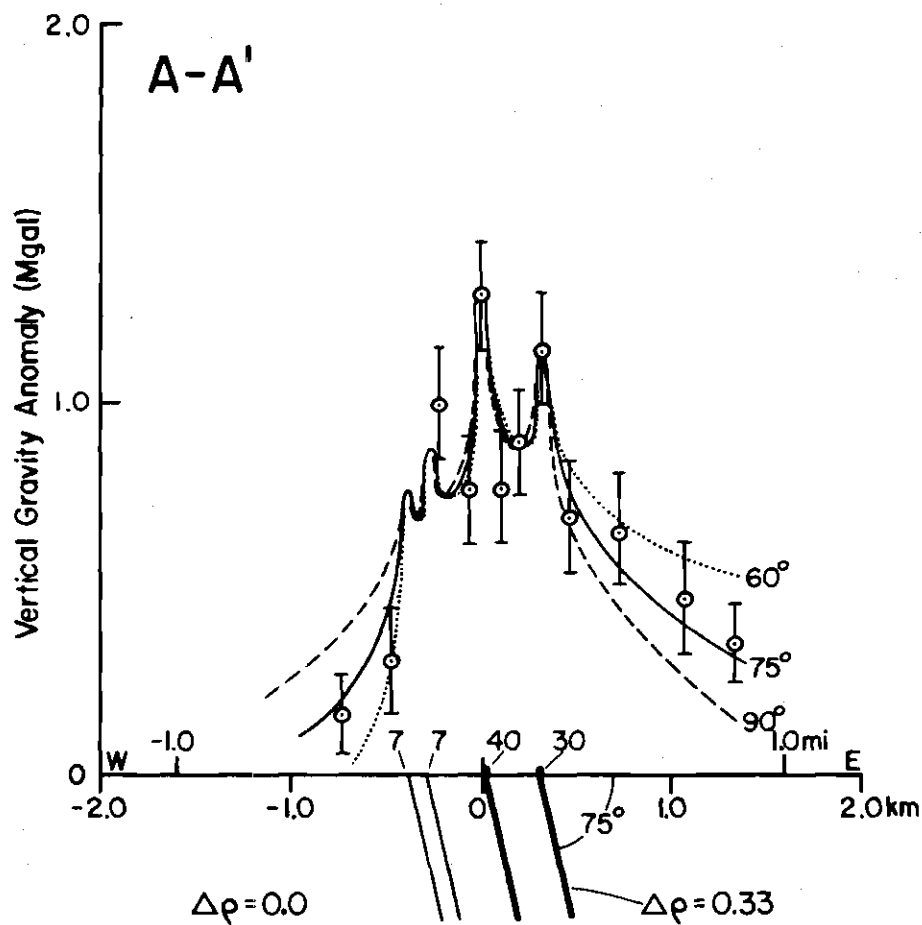


Figure 12. Detailed Gravity Profile A-A'. (Model dikes extend to infinite depth. Thickness of the modeled dike is given by the numeral above the dike. Heavy bar on profile indicates observed dike outcrop.)



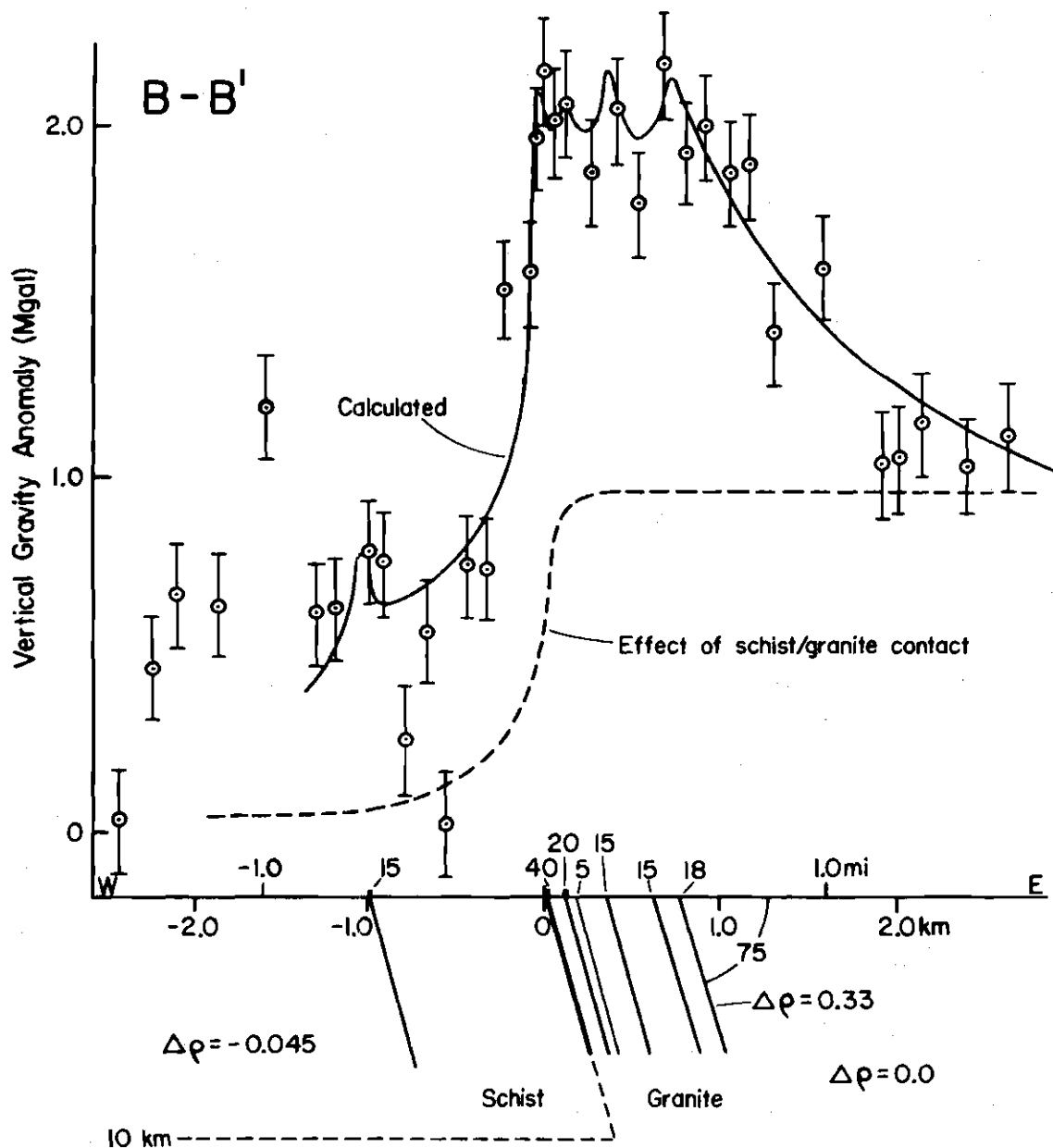


Figure 13. Detailed Gravity Profile B-B'. (Model dikes extend to infinite depth. Thickness of the modeled dike is given by the numeral above the dike. Heavy bar on profile indicates observed dike outcrop.)

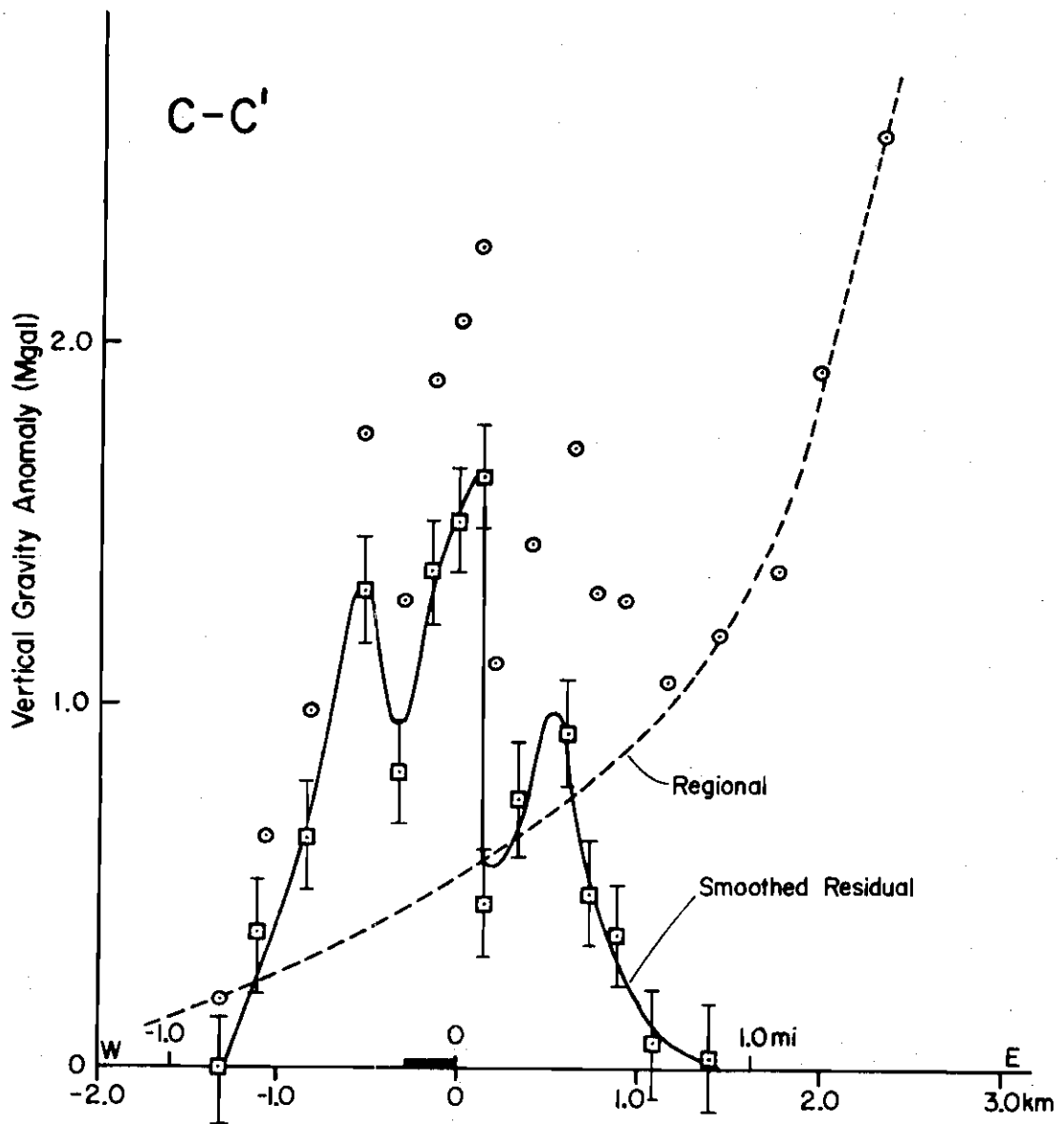


Figure 14. Detailed Gravity Profile C-C'. (Smoothed residual exhibits three definite peaks. Heavy bar on profile indicates observed dike outcrop.)

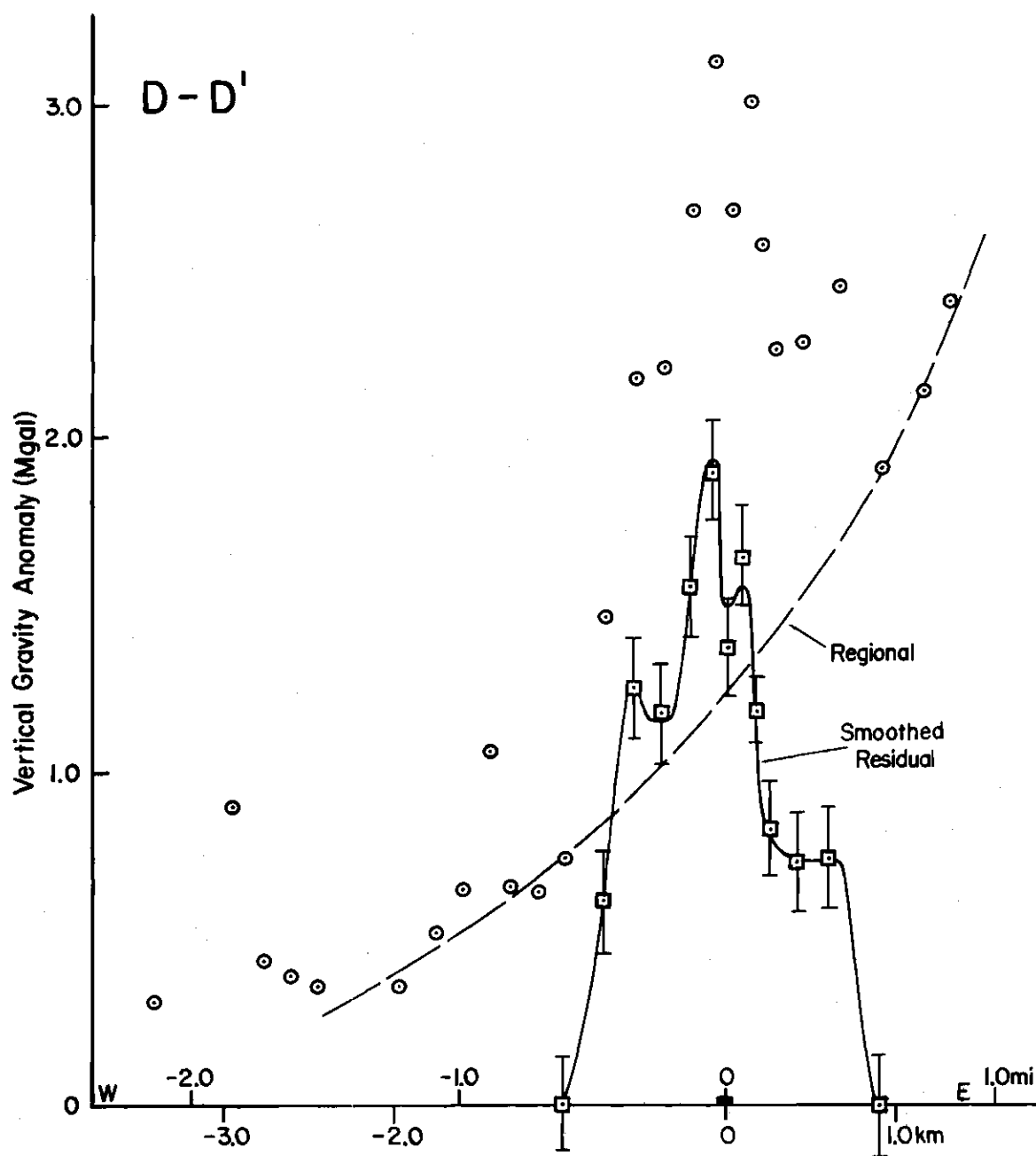


Figure 15. Detailed Gravity Profile D-D'. (Smoothed residual exhibits multiple peaks. Heavy bar on profile indicates observed dike outcrop.)

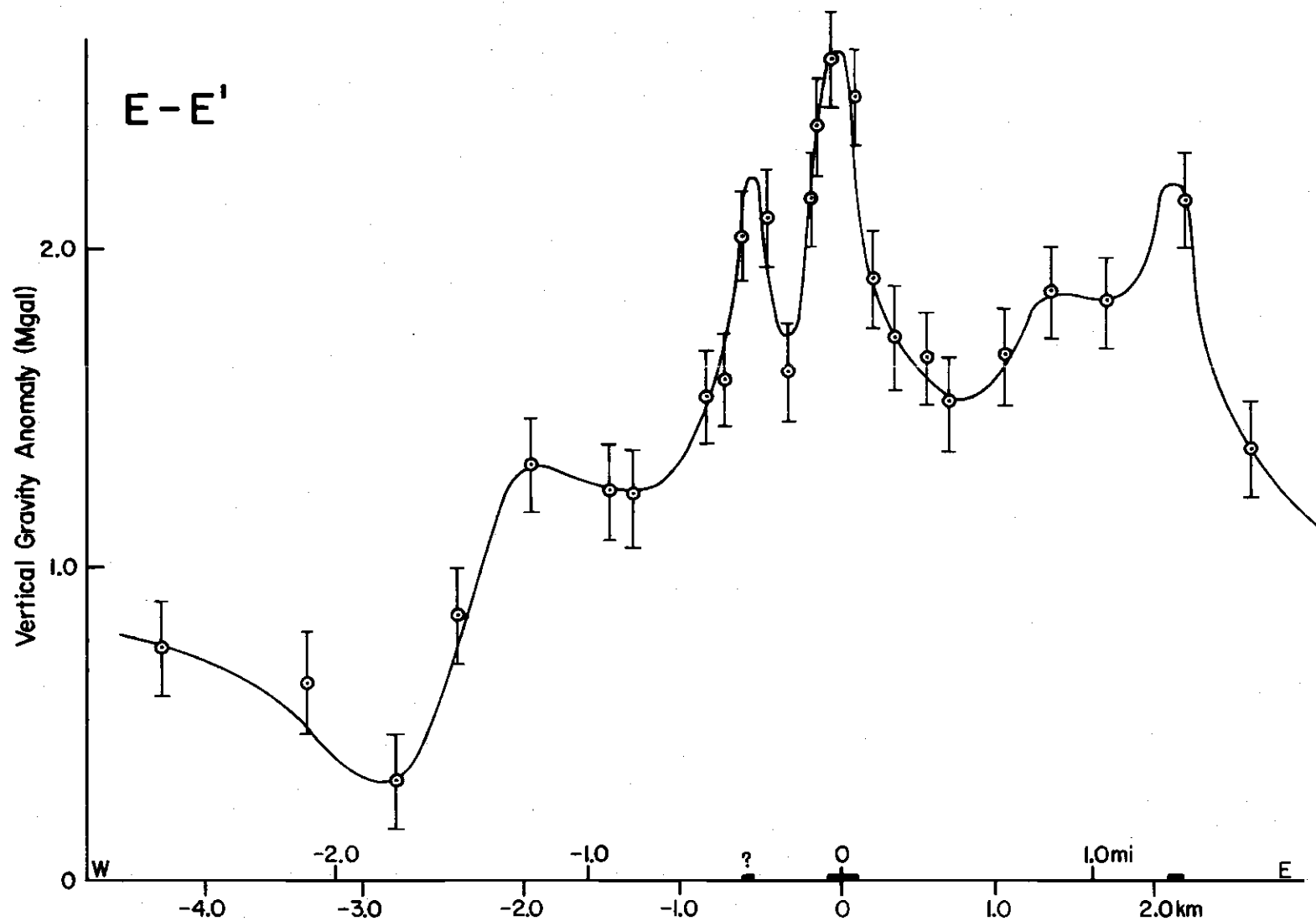


Figure 16. Detailed Gravity Profile E-E'. (Georgia Highway 16, south of Newnan. Heavy bar on profile indicates observed dike outcrops.)

gravity anomaly (Figure 12) consists of three peaks of about one milligal amplitude with the side and less prominent peaks spaced about 0.2 kilometers to either side of the main peak. Because there exist no known density anomalies, with exception of the diabase, along this profile and because the regional isogals are slightly curved in the region of this profile (Figure 4), no regional trend was removed. The three peaks and the asymmetry of the anomaly suggest that three dikes dipping to the east are necessary to account for the observed anomaly. Examination of the ground-level magnetic data for the same profile (Figure 8) suggests that the side dike to the west is actually two thin dikes. Assuming the existence of the four dikes, anomalies were computed using the observed outcrop widths for the main dike (40 meters) and the east dike (30 meters). The west dikes were not found to outcrop and because of the lesser magnitude of the associated gravity anomaly as compared to those of the main dike, they were modeled to be 7.5 meters wide.

Examination of the computed anomalies for dips of  $90^\circ$ ,  $75^\circ$  and  $60^\circ$  (Figure 12) shows the observed anomaly to fall between the computed curves for  $60^\circ$  and  $75^\circ$  with a dip angle of  $70^\circ$  (by interpolation) probably best satisfying the observed data.

#### Profile B-B'.

Detailed gravity line B-B' (Figure 4) was established along Georgia State Highway 109 between Greenville and Gay, Georgia. The observed anomaly consists of a "noisy" asymmetric broad peak (Figure 13). Amphibolites were found outcropping to the west of the main dike and could be responsible for the observed scatter in the data west of the dike. Diabase float was also found 1.0 kilometers west of the main

dike indicating that part of the observed scatter in the data may be a result of a side dike. A corresponding magnetic anomaly was encountered during the ground-level magnetic survey of the same profile (Figure 8).

Aside from the scatter in the data to the west of the main dike, there is a difference in the base level of the observed gravity. The east side of the dike is about 1.0 milligal more positive than the west side. This offset can be accounted for by the contact (no fault is visible and the orientation of the contact is unknown) between the Loachapoka Schist to the west and the granite to the east.

After the effect of the contact was removed, the resulting asymmetric residual anomaly could be modeled by a set of six dikes dipping to the east. Only the main dike was observed to outcrop, but the ground-level magnetic data (Figure 8) indicate the existence of the others. Determination of the angle of dip was not attempted because of the uncertainty in the location and orientation of the schist-granite contact.

#### Profile C-C'.

Profile C-C' (Figure 4) was established along the dirt road about 1.5 kilometers south of Georgia State Highway 109. The strong eastward positive trend of the observed data (Figure 14) is probably caused by granitic rocks which dip westward (Figure 3). Because the exact location and attitude of the granite-schist contact were unknown, a regional trend was estimated (Figure 14). The residual anomaly (Figure 14) consists of a central main peak and two secondary peaks on either side of the main peak. Although the dike was found to outcrop only coincident with the central peak of the gravity anomaly, the secondary peaks of the anomaly are probably due to side dikes. A similar phenomenon was

noted previously for profiles A-A' and B-B' further to the north. Because of the ambiguity created by the removal of the regional gravity trend by smoothing, and the noisy character of the ground-level magnetic data taken for the same profile (Figure 8), no attempt was made to determine the dip angle of the dikes.

#### Profile D-D'.

Line D-D' (Figure 4) was established along the paved county road about 8.0 kilometers southeast of Greenville and 4.5 kilometers south of detailed gravity line C-C'. As in profile C-C', there is a strong positive regional trend to the east (Figure 15), which was removed. The multiple peaks in the residual anomaly, (Figure 15) although not as well separated as in the previous profiles, again indicates that more than one dike is responsible for the observed anomaly. The only dike found to outcrop was again coincident with the main peak.

#### Profile E-E'.

Gravity line E-E' lies along Georgia Highway 16, southeast of Newnan. This profile, which is outside the main area of investigation, was obtained for comparison to the four previous profiles which lie within the area of detailed study 20 kilometers to the south. The observed anomaly (Figure 16) consists of three peaks all coincident with observed outcrops of dikes with the central peak corresponding to the widest outcrop. Whereas the side dikes in profiles to the south are located no more than 1.0 kilometers to the side of the main dike, the east dike on profile E-E' is located 2.1 kilometers from the main dike.

## CHAPTER V

### GROUND-LEVEL MAGNETICS AND THE ROLE OF NATURAL REMANENT MAGNETIZATION IN THE OBSERVED ANOMALIES

#### Ground-Level Magnetism

Four detailed ground-level magnetic (total field) profiles were obtained in order to examine the structure of the Meriwether dike and to provide additional data for the purpose of recontouring the aeromagnetic data previously described. Three of the profiles were obtained along the same roads as detailed gravity profiles A-A', B-B', and C-C'. The fourth was a closure profile connecting A-A' and B-B' (Figure 7). The data was obtained and reduced by standard techniques and are listed in Appendix II.

The reduced data (Figure 8) showed the main dike to be represented by an asymmetric anomaly for all three profiles. However, the calculated anomaly due to induction magnetization for a dike dipping  $70^{\circ}$  toward  $N70^{\circ}E$  is a symmetric positive peak (Figure 11). This suggests that induction magnetization is not the only cause for the observed anomaly and that natural remanent magnetization (NRM) must also be considered in the analysis of magnetic anomalies (Hood, 1963).

#### Natural Remanent Magnetization of the Meriwether Dike

Samples from a single outcrop of the Meriwether dike were collected and analyzed for NRM by Doyle Watts (personal communications). Intensities and directions for the NRM's are given in Appendix IV. A



Schmidt stereographic projection of the NRM directions (Figure 17) shows the NRM directions to be different from today's magnetic field. Samples from the center of the dike constitute one set of directions and those from the chilled margins another set. Because it is not known what proportion of the dike has which magnetization, an average of the directions (Figure 17) and magnitudes were used for computing the anomaly due to remanent magnetization. The average direction has a declination of  $28^{\circ}$  and an inclination of  $30^{\circ}$ . The average magnitude of the NRM is 0.00175 cgs.

#### Theoretical Anomalies and the Determination of the Koenigsberger Ratio, Q

The ratio of remanent magnetization to induced magnetization, the Koenigsberger ratio, Q is commonly cited in paleomagnetic studies as an indication of whether or not samples have been subjected to lightning strikes. It has more recently been construed (Green, 1960) as a measure of the importance of remanent magnetization in the analysis of magnetic anomalies. Using the calculated bulk susceptibility for the diabase of the Meriwether dike and the average intensity of NRM, the probable range of Q is 0.25 to 0.50. The range of Q's is a result of the uncertainty in the bulk susceptibility.

To examine the effects of Q on the observed anomalies, several models were computed using the method of Talwani and Heirtzler (1965). For a Q of 0.25 and bulk susceptibility of 0.013 cgs, the remanent magnetization causes the anomaly to become slightly asymmetric (Figure 18) with a peak to trough ratio of about 15 to 1. For bulk suscepti-

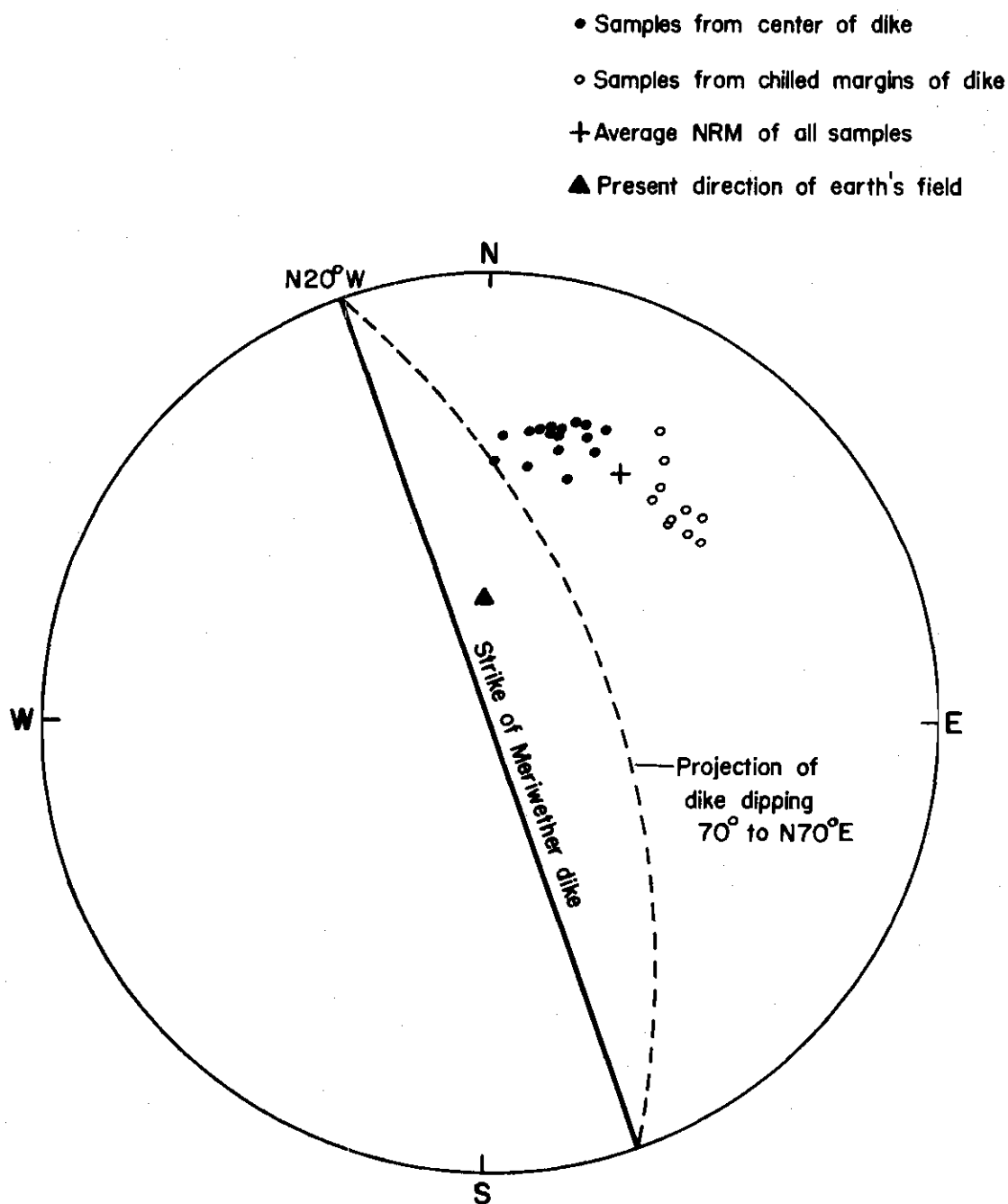


Figure 17. Stereographic Projection (Schmidt net) of Measured Natural Remanent Moments (NRM) for the Meriwether Dike.

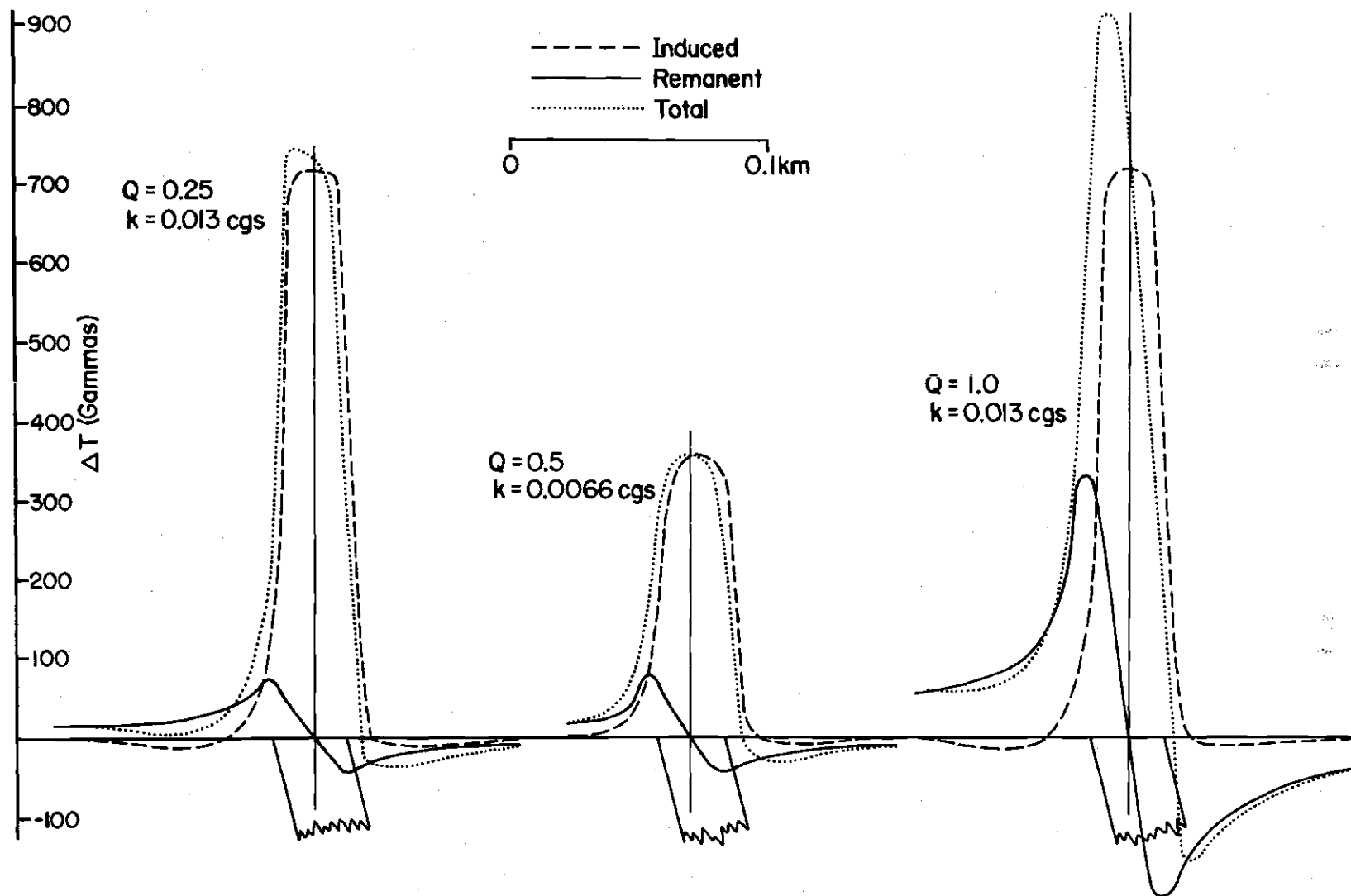


Figure 18. Calculated Curves Showing the Effect of Natural Remanent Magnetism on the Total Field Anomaly. (Dike is 30 meters wide and dips at  $75^\circ$ .)

bility 0.0066 cgs and corresponding  $Q$  of 0.5 the anomaly is found to be more asymmetric with a peak to trough ratio of 10 to 1.

The observed anomalies (Figure 8) however, exhibit an average peak to trough ratio of only 6. Further, the amplitude of the observed anomalies is greater than that calculated for the given range of susceptibilities and corresponding magnetizations. An increase in  $Q$  to 1.0 at a susceptibility of 0.013 cgs produced an asymmetric anomaly with a peak to trough ratio of about 6 to 1, (Figure 18) indicating that the bulk remanent magnetization was probably greater than that suggested by the surface samples.

Strangway (1965), who sampled a diabase dike at both the surface and at depth in a mine, found the ratio of remanent to induced magnetization to be greater for the underground samples. As a possible cause of this phenomenon Strangway (1965) suggests that temperature fluctuation at the surface, which was probably exposed for a considerable length of time, has accelerated the decay of the remanent magnetization. The same type of process may have occurred in the Meriwether dike and hence, the effective  $Q$ , which includes the effect of sub-surface portions of the dike, is probably closer to 1.0.

It should be noted that even if the dike actually dips as much as  $80^\circ$ , the magnetic anomaly due to induction would still be a symmetric peak (Figure 11), and hence the calculated  $Q$  of 1.0 would not be seriously affected.

## CHAPTER VI

## DISCUSSION AND CONCLUSIONS

Examination of the Meriwether dike by means of detailed gravity and magnetic profiles shows the dike to consist of an injection zone of dikes. Because the gravity anomalies of the component dikes are only partially separated, determination of whether or not the side dikes were branches of the main dike was not possible. On the basis of theoretical curves for the dike system dipping at various angles, the Meriwether dike (system) was estimated to dip at  $70^{\circ}$  toward  $N70^{\circ}E$ .

The  $70^{\circ}$  dip is not as steep as noted by Lester and Allen (1950), who found several of the larger dikes in Georgia to have a constant dip toward the east of  $75$  to  $90^{\circ}$  and Privett (1966) who found the diabase dikes in central South Carolina to dip  $80$  to  $90^{\circ}$  to the NE. It is possible, however, that the Meriwether dike may actually have a dip greater than  $70^{\circ}$ , but due to the ambiguity of the regional trend it is not possible to resolve how much greater.

A simple Bouguer gravity map compiled for central Meriwether County shows the Meriwether dike to be responsible for an anomaly of one to two milligals in the regional trend. After recontouring of the available aeromagnetic data, the Meriwether dike(s) proved to be the most prominent magnetic feature of the area, and this suggests that the sharpness of the anomaly and the flight line spacing suppressed the dike(s) in previous contouring. It is possible that in other areas of

eastern North America covered by aeromagnetic maps, other occurrences of diabase dikes may be similarly suppressed, thus reducing the probability of locating such dikes by their magnetic anomalies in areas not yet geologically mapped. Locating other dikes by this method may also be hampered by the effects of remanent magnetization.

Although not studied in detail, examination of high altitude infrared photographs (N.A.S.A., 1970) of central Meriwether County, revealed the existence of a linear anomaly coinciding with the Meriwether dike at known outcroppings in the northern third of the study area. The anomaly is probably due to a change of intensity of reflected infrared radiation from the vegetation growing in the soil derived from the diabase. It is suggested that this technique might be applied to other areas of eastern North America to determine outcrop patterns of the dikes in areas not yet geologically mapped.

Many problems currently exist concerning the diabase dikes which surround the North Atlantic Ocean such as dating, distribution and chemistry. It is hoped that more economic and efficient methods such as those suggested in this study will lead to the location of unmapped dikes, thus presenting a more complete picture of the distribution and role of these dikes in the history of the opening of the North Atlantic Ocean.

## APPENDICES

## APPENDIX I

## GRAVITY SURVEYS AND DATA REDUCTION

Gravity data were collected and reduced by standard techniques (Dobrin, 1960). The data consist of 242 points in total including base stations - 201 gravity observations are along five detailed lines and 41 gravity observations are regional data.

The gravimeter used was the Worden Educator, Model 113. The reading precision of this meter, when read by a single operator, is approximately  $\pm 0.1$  milligals. Instrumental drift for the Worden gravimeter for an eight hour period is typically 0.2 milligals, but can be as high as 0.5 milligals for the same period. The gravimeter was stored in the field vehicle each night preceding a survey to minimize the instrumental drift due to temperature change. The uncertainty in drift and reading precision combined give a gravity precision of  $\pm 0.2$  milligals. In all, three different base stations (Dorman and Ziegler, in preparation, Georgia Department of Mines, Mining and Geology) were used to correct for instrumental drift. Pertinent data concerning these base stations are given in Table 1.

The observed drifts are given in Table 2.



Table 1. Gravity Base Stations Used in this Study

Name	Base Number	Location	Gravity Value	Estimated Precision
Atlanta D	4	Georgia Tech	979527.37	$\pm 0.023$
LaGrange	33	City Hall	979484.42	$\pm 0.014$
Greenville	63	Meriwether County Court House	979489.58	$\pm 0.05$

Table 2. Instrumental Drift for Worden Gravimeter

Station Nos.	Georgia Tech Gravity Survey No.	Drift (Mgals/hr)	Time Between Base Stations
1 - 49	36	-0.074	6.1 hours
1 - 33	46	-0.018	9.8
1 - 48	51	0.000	7.4
48 - 72	51	-0.109	3.9
72 - 95	51	0.016	2.3
4 - 20	52	-0.053	2.8
20 - 41	52	-0.121	3.8

The gravity data were corrected for latitude effect by using the international gravity formula of 1930 (as given in Dobrin, 1960, page 187) and drift corrections for the meter by assuming linear meter drift between subsequent occupations of the base station. The standard Bouguer reduction density of  $2.67 \text{ gm/cm}^3$  was used to compute the Bouguer anomalies.

Elevation and location control were obtained from the following U. S. Geological Survey,  $7\frac{1}{2}$  minute topographic maps: Greenville (1971), Gay (1971), Warm Springs (1971), and Woodbury (1971). Where possible bench marks ( $\pm 1$  foot) or intersection elevations (assumed  $\pm 5$  feet) were used. Elevations of stations for which bench marks or intersections were not available were obtained by interpolating between contour lines (20 foot contour interval) and confirming those interpolations with barometric altimetry data. Elevations obtained by this technique were estimated to be plus or minus five feet. The resulting uncertainty in the reduced gravity values is thus  $\pm 0.35$  milligals. Because of the relatively shorter distance between stations along detail lines, and hence shorter times between gravity and barometric readings errors in drift for these measurements were considered less. This results in an estimated precision of  $\pm 0.2$  milligals for the reduced gravity values for stations along detail lines.

The data, in the standard Department of Defense computer card format, are listed by individual surveys in Table 3 (Figure 19).















## APPENDIX II

### GROUND LEVEL MAGNETIC SURVEYS AND DATA REDUCTION

Measurements of the magnitude of the geomagnetic field were made using a Geometrics Model G-816 proton-precession magnetometer. The digital display has a resolution of  $\pm 1$  gamma. The sensing element of the magnetometer is held at the end of an eight foot aluminum staff to suppress the magnetic effects of iron debris, e.g., beverage cans, small underground pipes, etc.

The data were reduced using standard techniques (Dobrin, 1960). The data consist of three detailed lines along graded dirt roads and a fourth along a two lane asphalt highway (Georgia State Highway 109). Approximately 0.016 kilometer was paced off between stations. Division of the actual length of the profile line by the number of station intervals yielded the actual station spacing. Times, needed for drift corrections, were recorded every ten minutes and the stations between time readings were assigned a time by assuming a linear sampling rate.

The variation in the main field over the area of investigation has been removed by approximating a gradient of 8.45 $\gamma$ /mile at N3 $^{\circ}$ W (Garland, 1971, Figure 17.2) by a simple latitude correction of 9.56 $\gamma$ /minute. Corrections for the diurnal variation were made using magnetograms of the Geomagnetic Observatory in Fredericksburg, Virginia (Figures 20 and 21) which were obtained from the World Data Center A of the National Oceanographic and Atmospheric Administration in

Boulder, Colorado.

These records were digitized and the vertical and horizontal components were added vectorily to give the magnitude of the total field. The difference between the computed total field and the value for the base line was taken to be the diurnal variation in the total field at Fredericksburg. Since the diurnal drift at any one place on the earth has been shown to be directly related to the hour angle of the sun (Matshusita and Campbell, 1965, Chapter 3) a shift of twenty-eight minutes was made in the drift curve so that it could be applied to the area of investigation. The variation of the diurnal drift with respect to latitude is negligible for the difference between Fredericksburg, Virginia and Greenville, Georgia (see Matshusita and Campbell, 1965, Figure 8, pages 321-323).

The computed latitude and longitude, time (Eastern Daylight Savings), raw magnetic value and corrected total magnetic field for each station occupied are given in Table 4.

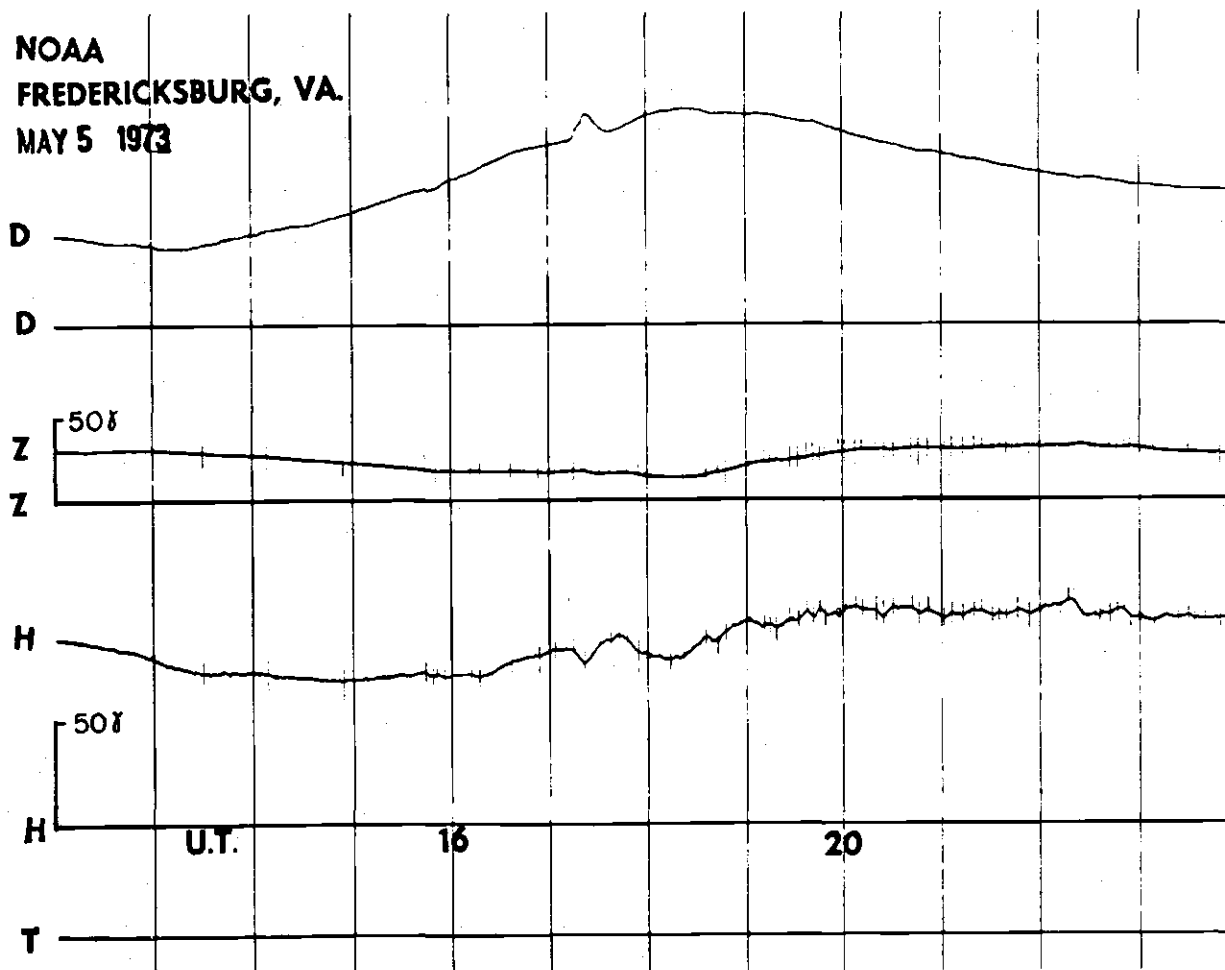


Figure 20. Magnetogram for May 5, 1973. (Courtesy National Oceanographic and Atmospheric Administration.)

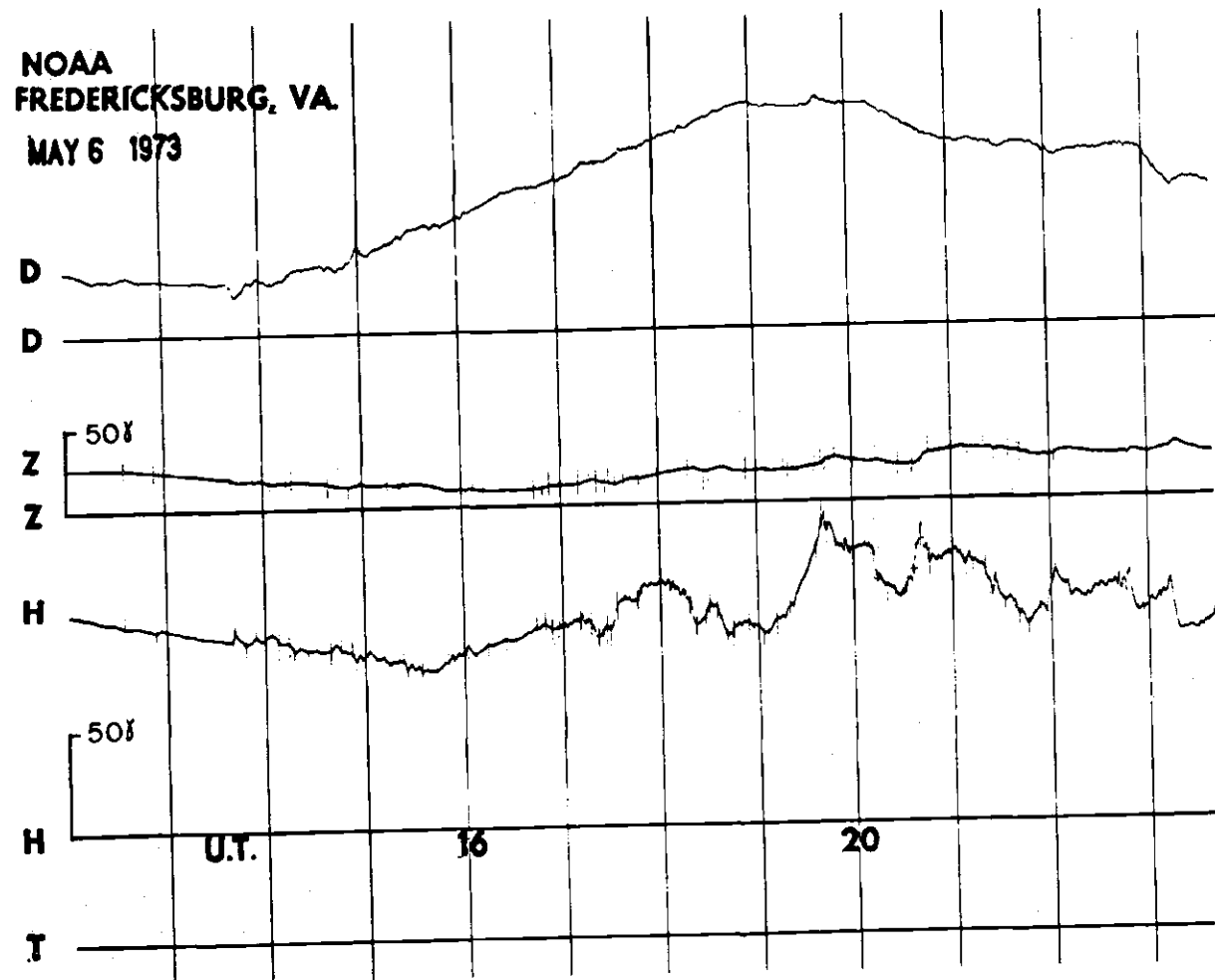


Figure 21. Magnetogram for May 6, 1973. (Courtesy National Oceanographic and Atmospheric Administration.)





Table 4. (Continued)

Profile A-A' (Concluded)

LATITUDE	LONGITUDE	TIME	RAW	REDUCED
1234567890123456789012345678901234567890				
33	4.875	84 39.190	18.680	53549 53581
33	4.875	84 39.170	18.700	53564 53596
33	4.875	84 39.150	18.713	53611 53644
33	4.875	84 39.130	18.725	53614 53647
33	4.875	84 39.110	18.738	53583 53616
33	4.875	84 39.090	18.750	53588 53621
33	4.875	84 39.070	18.763	53584 53617
33	4.875	84 39.050	18.776	53594 53627
33	4.875	84 39.030	18.788	53571 53603
33	4.875	84 39.010	18.801	53565 53597
33	4.875	84 38.990	18.813	53557 53589
33	4.875	84 38.980	18.820	53612 53644
33	4.875	84 38.970	18.826	53701 53732
33	4.875	84 38.960	18.832	53587 53618
33	4.875	84 38.940	18.845	53637 53668
33	4.875	84 38.930	18.851	53602 53633
33	4.875	84 38.920	18.857	53822 53853
33	4.875	84 38.910	18.864	53697 53728
33	4.875	84 38.890	18.876	53610 53640
33	4.875	84 38.870	18.889	53716 53746
33	4.875	84 38.850	18.901	53559 53589
33	4.875	84 38.830	18.914	53847 53877
33	4.875	84 38.820	18.920	53770 53800
33	4.875	84 38.810	18.927	53852 53882
33	4.875	84 38.800	18.933	53760 53790
33	4.875	84 38.780	18.946	53698 53728
33	4.875	84 38.760	18.958	53920 53950
33	4.875	84 38.740	18.971	53680 53710
33	4.875	84 38.720	18.983	53576 53605
33	4.875	84 38.700	18.994	53548 53577
33	4.875	84 38.680	19.006	53572 53601
33	4.875	84 38.660	19.017	53610 53639
33	4.875	84 38.640	19.028	53640 53669
33	4.875	84 38.620	19.039	53624 53653
33	4.875	84 38.600	19.050	53617 53646
33	4.875	84 38.580	19.061	53616 53645
33	4.875	84 38.560	19.072	53635 53664
33	4.875	84 38.540	19.083	53668 53697
33	4.875	84 38.520	19.094	53654 53683
33	4.875	84 38.500	19.106	53594 53623
33	4.875	84 38.480	19.117	53670 53699
33	4.875	84 38.460	19.131	53640 53669
33	4.875	84 38.440	19.144	53628 53657
33	4.875	84 38.420	19.158	53628 53657
33	4.875	84 38.400	19.172	53623 53652
33	4.875	84 38.380	19.186	53645 53674
33	4.875	84 38.360	19.200	53677 53706
33	4.875	84 38.340	19.212	53643 53672
33	4.875	84 38.320	19.224	53608 53637
33	4.875	84 38.300	19.236	53599 53628
33	4.875	84 38.280	19.248	53598 53627
33	4.875	84 38.260	19.260	53605 53634
33	4.875	84 38.240	19.271	53605 53634
33	4.875	84 38.220	19.283	53592 53621
1234567890123456789012345678901234567890				





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Table 4. (Continued)

Profile C-C', May 6, 1973

LATITUDE	LONGITUDE	TIME	RAW	REDUCED
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890
33	1.945	84 40.277	12.050	53593 53624
33	1.944	84 40.259	12.060	53595 53626
33	1.942	84 40.242	12.070	53561 53592
33	1.941	84 40.224	12.080	53548 53578
33	1.940	84 40.206	12.090	53561 53591
33	1.939	84 40.188	12.100	53709 53739
33	1.938	84 40.179	12.109	53535 53565
33	1.937	84 40.171	12.118	53594 53624
33	1.936	84 40.153	12.136	53701 53731
33	1.935	84 40.135	12.154	53620 53650
33	1.949	84 40.125	12.172	53609 53638
33	1.963	84 40.116	12.190	53705 53734
33	1.977	84 40.106	12.208	53595 53624
33	1.991	84 40.096	12.226	53659 53688
33	2.005	84 40.087	12.244	53655 53683
33	2.019	84 40.077	12.262	53669 53697
33	2.032	84 40.068	12.281	53655 53683
33	2.046	84 40.058	12.299	53661 53689
33	2.060	84 40.048	12.317	53661 53689
33	2.074	84 40.039	12.327	53657 53685
33	2.088	84 40.029	12.337	53627 53650
33	2.102	84 40.019	12.347	53637 53660
33	2.116	84 40.010	12.357	53572 53600
33	2.130	84 40.000	12.367	53535 53563
33	2.132	84 39.990	12.377	53537 53565
33	2.134	84 39.980	12.387	53541 53569
33	2.136	84 39.970	12.397	53542 53570
33	2.138	84 39.960	12.407	53554 53582
33	2.142	84 39.940	12.427	53573 53601
33	2.146	84 39.920	12.447	53554 53582
33	2.151	84 39.900	12.467	53615 53644
33	2.155	84 39.880	12.467	53624 53653
33	2.159	84 39.860	12.467	53558 53587
33	2.163	84 39.840	12.708	53614 53643
33	2.165	84 39.830	12.719	53631 53660
33	2.161	84 39.822	12.729	53609 53638
33	2.157	84 39.813	12.740	53588 53617
33	2.153	84 39.805	12.750	53599 53629
33	2.144	84 39.789	12.768	53582 53612
33	2.136	84 39.772	12.787	53585 53615
33	2.128	84 39.756	12.805	53580 53610
33	2.120	84 39.739	12.824	53563 53593
33	2.111	84 39.723	12.842	53590 53620
33	2.103	84 39.706	12.860	53540 53570
33	2.095	84 39.690	12.879	53552 53582
33	2.091	84 39.681	12.888	53545 53576
33	2.087	84 39.673	12.897	53529 53560
33	2.082	84 39.665	12.906	53520 53551
33	2.078	84 39.657	12.916	53533 53564
33	2.074	84 39.648	12.925	53605 53636
33	2.070	84 39.640	12.934	53741 53772
33	2.066	84 39.634	12.943	53826 53857
33	2.060	84 39.627	12.952	53765 53796
33	2.056	84 39.621	12.961	53667 53698
33	2.051	84 39.615	12.971	53944 53975
33	2.046	84 39.609	12.980	54188 54219
33	2.041	84 39.602	12.989	53668 53700
33	2.036	84 39.596	12.998	53621 53653
33	2.031	84 39.590	13.007	53727 53759
33	2.027	84 39.583	13.017	53609 53641
33	2.022	84 39.577	13.100	53643 53676
33	2.017	84 39.571	13.106	53613 53646
33	2.012	84 39.564	13.111	53624 53657
33	2.007	84 39.558	13.117	53618 53652
33	1.998	84 39.546	13.128	53626 53660
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890

LATITUDE	LONGITUDE	TIME	RAW	REDUCED
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890
33	1.988	84 39.533	13.139	53699 53733
33	1.979	84 39.520	13.150	53809 53843
33	1.974	84 39.514	13.156	53689 53723
33	1.964	84 39.501	13.167	53680 53715
33	1.954	84 39.489	13.180	53605 53640
33	1.950	84 39.483	13.187	53577 53612
33	1.940	84 39.470	13.200	53627 53662
33	1.940	84 39.452	13.213	53650 53685
33	1.939	84 39.433	13.227	53585 53620
33	1.939	84 39.424	13.233	53654 53689
33	1.939	84 39.415	13.242	53677 53712
33	1.939	84 39.406	13.250	53743 53778
33	1.939	84 39.397	13.258	53648 53683
33	1.939	84 39.388	13.267	53671 53706
33	1.939	84 39.378	13.275	53693 53729
33	1.938	84 39.369	13.283	53710 53746
33	1.938	84 39.360	13.292	53668 53704
33	1.938	84 39.351	13.300	53694 53730
33	1.938	84 39.342	13.308	53785 53821
33	1.938	84 39.332	13.317	53871 53907
33	1.938	84 39.323	13.325	53991 54027
33	1.938	84 39.314	13.333	53704 53740
33	1.938	84 39.305	13.342	53744 53780
33	1.937	84 39.296	13.350	53641 53677
33	1.937	84 39.287	13.358	53770 53806
33	1.937	84 39.278	13.367	53650 53686
33	1.937	84 39.268	13.375	53773 53809
33	1.937	84 39.259	13.383	53781 53817
33	1.937	84 39.250	13.392	53785 53821
33	1.937	84 39.241	13.400	53795 53831
33	1.936	84 39.232	13.408	53844 53880
33	1.936	84 39.223	13.417	53863 53900
33	1.936	84 39.213	13.427	53833 53870
33	1.936	84 39.204	13.438	53931 53968
33	1.936	84 39.195	13.448	53947 53984
33	1.936	84 39.186	13.459	54494 54531
33	1.936	84 39.177	13.470	54314 54351
33	1.935	84 39.168	13.480	54259 54296
33	1.935	84 39.158	13.491	53926 53963
33	1.935	84 39.149	13.502	53849 53886
33	1.935	84 39.140	13.512	53525 53562
33	1.935	84 39.131	13.523	53609 53646
33	1.935	84 39.122	13.533	53626 53663
33	1.935	84 39.113	13.544	53545 53583
33	1.934	84 39.103	13.555	53585 53623
33	1.934	84 39.094	13.565	53635 53673
33	1.934	84 39.085	13.576	53599 53637
33	1.934	84 39.076	13.586	53600 53638
33	1.934	84 39.067	13.597	53685 53723
33	1.934	84 39.048	13.618	53729 53767
33	1.933	84 39.030	13.639	53783 53821
33	1.933	84 39.021	13.650	53831 53870
33	1.933	84 39.012	13.659	53761 53800
33	1.933	84 39.003	13.668	53770 53809
33	1.933	84 38.984	13.686	53958 53997
33	1.933	84 38.975	13.695	53773 53812
33	1.932	84 38.966	13.705	53806 53845
33	1.932	84 38.957	13.714	53748 53787
33	1.932	84 38.938	13.732	53855 53895
33	1.932	84 38.920	13.750	53732 53771
33	1.931	84 38.902	13.761	53716 53755
33	1.931	84 38.883	13.773	53707 53745
33	1.931	84 38.865	13.784	53686 53724
33	1.931	84 38.847	13.796	53647 53684
33	1.930	84 38.828	13.807	53760 53797
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890

Table 4. (Continued)

Profile C-C' (Concluded)

LATITUDE	LONGITUDE	TIME	RAW	REDUCED
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890
33	1.930	84 38.810	13.819	53735 53772
33	1.925	84 38.790	13.830	53723 53760
33	1.921	84 38.770	13.842	53623 53661
33	1.916	84 38.750	13.853	53712 53750
33	1.913	84 38.740	13.859	53670 53708
33	1.909	84 38.720	13.871	53724 53762
33	1.904	84 38.700	13.882	53729 53767
33	1.899	84 38.680	13.894	53722 53760
33	1.894	84 38.660	13.905	53773 53811
33	1.892	84 38.650	13.911	53823 53861
33	1.890	84 38.640	13.917	53791 53829
33	1.885	84 38.620	13.937	53660 53698
33	1.889	84 38.602	13.957	53826 53864
33	1.890	84 38.593	13.967	53803 53842
33	1.892	84 38.584	13.977	53754 53793
33	1.894	84 38.575	13.987	53809 53849
33	1.896	84 38.566	13.997	53793 53833
33	1.899	84 38.548	14.017	53735 53776
33	1.903	84 38.530	14.027	53808 53849
33	1.906	84 38.512	14.038	53848 53889
33	1.910	84 38.494	14.049	53857 53899
33	1.913	84 38.476	14.060	53773 53815
33	1.917	84 38.458	14.071	53783 53825
33	1.920	84 38.440	14.081	53661 53703
33	1.920	84 38.421	14.092	53708 53751
33	1.920	84 38.403	14.103	53609 53652
33	1.919	84 38.384	14.114	53649 53692
33	1.919	84 38.365	14.125	53662 53705
33	1.919	84 38.347	14.135	53730 53773
33	1.919	84 38.328	14.146	53720 53764
33	1.918	84 38.309	14.157	53801 53845
33	1.918	84 38.291	14.168	53735 53779
33	1.918	84 38.272	14.178	53702 53746
33	1.918	84 38.253	14.189	53623 53667
33	1.918	84 38.235	14.200	53590 53634
33	1.917	84 38.216	14.210	53586 53631
33	1.917	84 38.197	14.221	53596 53642
33	1.917	84 38.179	14.231	53636 53682
33	1.917	84 38.160	14.241	53653 53700
33	1.916	84 38.141	14.251	53669 53716
33	1.916	84 38.123	14.262	53646 53693
33	1.916	84 38.113	14.267	53593 53640
33	1.916	84 38.095	14.277	53586 53633
33	1.916	84 38.076	14.287	53578 53625
33	1.915	84 38.057	14.297	53578 53625
33	1.915	84 38.039	14.308	53563 53610
33	1.915	84 38.020	14.318	53603 53650
33	1.915	84 38.010	14.323	53593 53640
33	1.915	84 37.991	14.333	53669 53717
33	1.915	84 37.972	14.344	53663 53711
33	1.914	84 37.953	14.355	53683 53731
33	1.914	84 37.934	14.365	53679 53727
33	1.914	84 37.915	14.376	53653 53701
33	1.914	84 37.886	14.392	53616 53664
33	1.914	84 37.867	14.403	53641 53689
33	1.914	84 37.847	14.413	53663 53712
33	1.913	84 37.828	14.424	53661 53710
33	1.913	84 37.809	14.435	53670 53719
33	1.913	84 37.790	14.445	53633 53682
33	1.913	84 37.771	14.456	53633 53682
33	1.913	84 37.751	14.467	53620 53669
33	1.913	84 37.732	14.477	53599 53648
33	1.912	84 37.713	14.488	53632 53681
33	1.912	84 37.694	14.499	53679 53729
33	1.912	84 37.675	14.509	53626 53676
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890

LATITUDE	LONGITUDE	TIME	RAW	REDUCED
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890
33	1.912	84 37.656	14.520	53638 53688
33	1.912	84 37.636	14.531	53710 53760
33	1.912	84 37.617	14.541	53676 53726
33	1.911	84 37.598	14.552	53779 53829
33	1.911	84 37.579	14.563	53668 53718
33	1.911	84 37.560	14.573	53651 53701
33	1.911	84 37.540	14.584	53619 53669
33	1.911	84 37.521	14.595	53614 53663
33	1.911	84 37.512	14.600	53614 53663
33	1.910	84 37.493	14.608	53609 53658
33	1.910	84 37.473	14.617	53618 53667
33	1.910	84 37.454	14.625	53614 53663
33	1.910	84 37.435	14.633	53567 53616
1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890	1234567890123456789012345678901234567890



## APPENDIX III

## COMPUTER PROGRAMS FOR GRAVITY AND MAGNETICS MODELING

A computer program was developed for computing the vertical gravity anomaly caused by a hypothetical two-dimensioned structure using the method of Talwani, Worzel, and Landisman (1959). A listing of the main program and subroutines necessary for line printer plotting is given in Table 5.

A computer program was developed for computing horizontal, vertical, and total magnetic anomalies due to induction, NRM or mixed magnetization using the method of Talwani and Heirtzler (1965). A listing of the main program and all referenced subroutines is given in Table 6.

Table 5. Two-Dimensional Gravity Modeling Program

```

1234567890123456789012345678901234567890123456789012345678901234567890
C GRAVITY PROFILING FOR 2-DIMENSIONAL STRUCTURES
C GRAVITY FOR 2-DIMENSIONAL STRUCTURES(AFTER-TALWANI,WORTZEL, LANDISMANN)
C INPUT(2I10,3F10.3) CARD NO. ONE
C LL=NO. OF POLYGONS, NX=NO. OF GRAVITY VALUES, DX=SEPARATION OF
C GRAVITY VALUES, XO=POSITION OF FIRST GRAVITY VALUE, SCALE=PLOT
C SCALE=FOR DRAW
C IF SCALE=0. PROGRAM CALCULATES SCALE
C IF SCALE=1. NO GRAPH IS DRAWN
C INPUT(FREE FIELD) LL CARDS
C NX=NO. OF CORNERS OF POLYGON TAKEN CLOCKWISE, DRHO=DENSITY
C CONTRAST, X(I,J),Z(I,J)=COORDINATES OF CORNERS-JTH CORNER OF ITH
C POLYGON
C REPEAT SEQUENCE FOR ADDITIONAL PROFILES
C BLANK CARD AT END TO TERMINATE CALCULATION
DIMENSION DRHO(50),X(50,20),Z(50,20),NN(50),XX(20),ZZ(20),GAL(500)
DIMENSION A(80)
25 READ (5,900)A
900 FORMAT(80A1)
READ (5,500) LL,NDX,DX,XO , SCALE
WRITE(6,901)A
901 FORMAT(1H1,80A1,/)
WRITE(6,503) LL,NDX,DX,XO
503 FORMAT(1H ,28HVERTICAL GRAVITY ANOMALY FOR 15, 9H POLYGONS/2X,
13HTHE 15,16H-GRAVITY VALUES,,F10.4,19H-KM.APART. BEGIN AT,F10.4)
IF(LL) 26,26,27
27 DO 100 I=1,LL
500 FORMAT (2I10,3F10.5)
READ(5,501) NXZ,DRHO(I), (X(I,J),Z(I,J),J=1,NX)
WRITE(6,502) NXZ,DRHO(I), (X(I,J),Z(I,J),J=1,NXZ)
501 FORMAT ( )
502 FORMAT(16H NO OF POINTS = ,I10,22H DENSITY DIFFERENCE = /,(1X,1,F10
<.3/))
100 NN(I) = NXZ
DO 101 I=1,LL
NNI = NN(I)
DO 101 J = 1,NNI
101 X(I,J) = X(I,J) -XO
DO 102 I=1,NDX
G=0.0
DO 103 J=1,LL
NNJ = NN(J)
DO 104 K = 1,NNJ
XX(K) = X(J,K)
104 ZZ(K)=Z(J,K)
CALL TWLZ(NN(J),XX,ZZ,DRHO(J),GA)
103 G= G+GA
GAL(I) = G
DO 110 M=1,LL
NNM = NN(M)
DO 110 N = 1, NNM
110 X(M,N) = X(M,N)-DX
102 CONTINUE
WRITE(6,505) (GAL(I),I=1,NX)
505 FORMAT(1X//24H GRAVITY ANOMALY IN MGAL,/(5F15.4))
ISC=SCALE
IF(ISC.EQ.13) GO TO 25
IF(SCALE) 130,131,130
131 CALL MXSCL(NDX,GAL,SCALE)
130 CALL DRAW(NDX,1,GAL,SCALE)
GO TO 25
26 STOP
END
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 5. (Concluded)

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1234567890123456789012345678901234567890123456789012345678901234567890
SUBROUTINE TWLZ(K1,XX,ZZ,DRHO,GA)
C USES METHOD OF TALWANI, WENZEL, AND LANDISMAN (JGR 1959 PP 49-59)
C TO GIVE GRAVITY ANOMALY AT X=0,Z=0. IN MGAL FOR TWO DIMENSIONED
C BODY IN VERTICAL PLANE DESCRIBED BY A POLYGON (IN KILOMETERS)
DIMENSION XX(K1),ZZ(K1)
PI = 3.141592654
KK = K1-1
GA = 0.0
DO 100 K=1,KK
  K2 = K+1
  IF(XX(K)*ZZ(K2)-XX(K2)*ZZ(K)) 30,100,30
30 IF(XX(K)-XX(K2)) 85,20,85
20 XZ = ((XX(K2)**2 + ZZ(K2)**2)/(XX(K)**2 + ZZ(K)**2))
  DG = 0.5*LOG(XZ)*XX(K)
  GO TO 99
85 IF(ZZ(K)-ZZ(K2)) 235,72,235
72 DG = ZZ(K)*(ATAN2(ZZ(K2),XX(K2))-ATAN2(ZZ(K),XX(K)))
  GO TO 99
235 A = (XX(K2)-XX(K))/(ZZ(K2)-ZZ(K))
  B = (XX(K)*ZZ(K2) - XX(K2)*ZZ(K))/(ZZ(K2)-ZZ(K))
  IF(XX(K)) 200,201,200
201 DG = (B/(1.+A*A))*(.5*LOG((XX(K2)*XX(K2)+ZZ(K2)*ZZ(K2))/(ZZ(K)*
  1 ZZ(K))) - A*(ATAN2(ZZ(K2),XX(K2))-PI/2.))
200 IF(XX(K2)) 31,210,31
210 DG = (A/(1.+A*A))*(.5*LOG((ZZ(K2)*ZZ(K2)/(XX(K)*XX(K)+
  1 ZZ(K)*ZZ(K))) + A*(ATAN2(ZZ(K),XX(K)) - PI/2.))
31 DG=LOG((XX(K2)*XX(K2)+ZZ(K2)*ZZ(K2))/(XX(K)*XX(K)+ZZ(K)*ZZ(K)))
  DG=(B/(1.0+A*A))*(0.5*DG-A*(ATAN2(ZZ(K2),XX(K2))-ATAN2(ZZ(K),
  1 XX(K))))
99 GA=(13.34) *DRHO*DG + GA
100 CONTINUE
RETURN
END

SUBROUTINE MXSCL(N,A,AMAX)
DIMENSION A(N)
AMAX = 0
DO 26 I = 1,N
  IF (ABS(A(I))-AMAX) 26,26,25
25 AMAX = ABS(A(I))
26 CONTINUE
AN = LOG10(AMAX)
IF(AN) 17,18,19
17 NN = AN - 1
  GO TO 20
19 NN = AN
20 IA = AMAX/(10.**NN)
  IF(IA.LE.2) GO TO 14
  IF(IA.LE.5) GO TO 15
16 AMAX = 10.*(10.**NN)
18 RETURN
14 AMAX = 2.*(10.**NN)
  RETURN
15 AMAX = 5.*(10.**NN)
  RETURN
END

SUBROUTINE DRAW (NTOT, INC, F, SCALE)
C NTOT=TOTAL NUMBER OF POINTS IN F. F IS THE DATA (ONE DIMENSIONAL)
C TO BE PLOTTED. INC IS THE SAMPLE INTERVAL FOR PLOTTING F.
C SCALE IS THE AMPLITUDE OF ONE FULL SCALE DEFLECTION
DIMENSION F(NTOT)
DATA AA1/1H /,AA2/1H /,AA3/1H /
WRITE(5,1011) SCALE,(I,I=-1,10) * (AA2,M=1,21)
1011 FORMAT(1H1,14.8,17H-MAGNALS FULL SCALE/3X,2015/2X,22A5)
10 DO 1501 K = 1, NTOT, INC
  FK = 50.*F(K)/SCALE
  KI = FK/50
  KK = FK - KI*50.+50.5
  WRITE (6,511) AA2, (AA1,I=1,KK),AA2
511 FORMAT (1X,110A1)
1501 CONTINUE
RETURN
END
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. Two-Dimensional Magnetics Modeling Program

```

123456789012345678901234567890123456789012345678901234567890
C  COMPUTATION OF MAGNETIC ANOMALIES CAUSED BY THE MAGNETIZATION OF TWO
C  DIMENSIONAL, (INFINITE EXTENT IN THIRD DIRECTION).
C  THIS PROGRAM IS A MODIFICATION OF THE PROGRAM BY
C  MANIK TALWANI AND JAMES R. HEIRTZLER WHICH WAS PUBLISHED IN
C  COMPUTERS IN THE MINERAL INDUSTRIES, PART 1, STANFORD UNIVERSITY
C  PUBLICATION GEOL. SCI., VOL 9, NO. 1, PAGES 464-480.
C  THIS PROGRAM BY GEORGE RUTH
C  SCHOOL OF GEOPHYSICAL SCIENCES
C  GEORGIA INSTITUTE OF TECHNOLOGY
C  ATLANTA, GA 30332
C  SEPTEMBER 20, 1972
      DIMENSION FLOPT(200),PCUM(200),QSUM(200),VANNLY(200,10,2),
      * HANNLY(200,10,2),TANNLY(200,10,2),SUS(10),REMEC(10),REMDIP(10),
      * REMMAG(10),VIND(200),VREN(200),HIND(200),HREM(200),
      * TIND(200),TREM(200),VTOT(200),HTOT(200),TTOT(200)
      DIMENSION Y(10,20),Z(10,20),VWCRK(200),HWORK(200),TWORK(200)
      DIMENSION WORK1(200),WORK2(200)
      REAL MOTION
      INTEGER PRIND,PLINDV,PLINDH,PLINDT,VCOMPI,VCOMPR,HCOMPI,
      * HCOMPR,TCOMPI,TCOMPR,PRNTAL,PLVAL,PLHAL,PLTAL
C*****
C***READ RUN DESCRIPTION CARD
      1 READ (5,3000) NPOLY,PRIND,PLINDV,PLINDH,PLINDT,VCOMPI,VCOMPR,
      * HCOMPI,HCOMPR,TCOMPI,TCOMPR,PRNTAL,PLVAL,PLHAL,PLTAL
      3000 FORMAT (15I2)
      IF (NPOLY.EQ.0) GO TO 520
C*****
C***NPOLY = NUMBER OF POLYGONS
C***THE FOLLOWING ARE OUTPUT OPTION VARIABLES
C  IF VARIABLE = 0, OPTION NOT EXERCISED
C  IF VARIABLE = 1, OPTION IS EXERCISED
C***PRIND = LIST ANOMALIES FOR EACH POLYGON
C***PLINDV = PLOT VERTICAL ANOMALY FOR EACH POLYGON
C***PLINDH = PLOT HORIZONTAL ANOMALY FOR EACH POLYGON
C***PLINDT = PLOT TOTAL ANOMALY FOR EACH POLYGON
C***VCOMPI = COMPOSITE PLOT OF VERTICAL ANOMALY DUE TO INDUCTION ONLY
C***VCOMPR = COMPOSITE PLOT OF VERTICAL ANOMALY DUE TO REMANENCE ONLY
C***HCOMPI = COMPOSITE PLOT OF HORIZONTAL ANOMALY DUE TO INDUCTION ONLY
C***HCOMPR = COMPOSITE PLOT OF HORIZONTAL ANOMALY DUE TO REMANENCE ONLY
C***TCOMPI = COMPOSITE PLOT OF TOTAL ANOMALY DUE TO INDUCTION ONLY
C***TCOMPR = COMPOSITE PLOT OF TOTAL ANOMALY DUE TO REMANENCE ONLY
C***PRNTAL = LINE PRINTER LISTING OF ANOMALIES DUE TO ALL POLYGONS
C***PLVAL = PLOT VERTICAL ANOMALY DUE TO ALL POLYGONS - REMANENT + INDUCED
C***PLHAL = PLOT HORIZONTAL ANOMALY DUE TO ALL POLYGONS - REMANENT + INDUCED
C***PLTAL = PLOT TOTAL ANOMALY DUE TO ALL POLYGONS - REMANENT + INDUCED
C*****
C***READ PROFILE DESCRIPTION CARD
C*****
      4 READ (5,1000) NFPTS,FRSTX,DELX,DECPLX,OBVSHT,HATCHM
      1000 FORMAT (13,5F8.3)
C*****
C***FRSTX = X COORDINATE OF THE FIELD POINT FARTHEST IN THE NEGATIVE
C***          DIRECTION(LEFT).
C***DELX = INTERVAL BETWEEN FIELD POINTS.
C***NFPTS = NUMBER OF FIELD POINTS
C***DECPLX = DECLINATION OF THE POSITIVE X AXIS OF THE PROFILE LINE,
C***          MEASURED POSITIVE CLOCKWISE FROM GEOGRAPHIC NORTH IN DEGREES.
C***OBVSHT = HEIGHT ABOVE PROFILE LINE AT WHICH ANOMALY WILL BE
C***          CALCULATED, MEASURED POSITIVE UP(SEE FOLLOWING FIGURE).
C***HATCHM = INTERVAL BETWEEN HATCHMARKS FOR PLOTS -- MUST BE AN
C***          INTEGRAL MULTIPLE OF DELX
123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
C*** FOR THIS POLYGON.
      CALL POLYPO (NFPTS, NSIDES, FLDPT, OBSVHT, PSUM, QSUM, I, X, Z, DELX)
C*** CALCULATE ANOMALY DUE TO INDUCTION FOR THIS POLYGON AT EACH FIELD POINT.***
      MGTON = SUS(I) * FLDPT
      DO 40 J = 1, NFPTS
        VANMLY(J,I,1) = -2.*MGTON*((CDIPR*CDIFFR*QSUM(J))-(SDIPR*PSUM(J)))
        HANMLY(J,I,1) = -2.*MGTON*((CDIPR*CDIFFR*PSUM(J))+(SDIPR*QSUM(J)))
C*** FOR ANOMALIES SMALL IN COMPARISON TO THE EARTH'S FIELD, THE INTENSITY OF THE
C*** TOTAL ANOMALY DUE TO INDUCTION IS THE SUM OF THE PROJECTIONS OF THE
C*** VERTICAL AND HORIZONTAL ANOMALIES DUE TO INDUCTION ALONG THE DIRECTION OF
C*** THE EARTH'S FIELD.
        TANMLY(J,I,1) = HANMLY(J,I,1)*CDIPR*CDIFFR + VANMLY(J,I,1)*SDIPR
      40 CONTINUE
C*** CALCULATE ANOMALY DUE TO REMANENT MAGNETIZATION FOR THIS POLYGON
444      40 8ARH 8IA=D POINH
        CDIP = COS (.0174533*REMDIP(I))
        SDIP = SIN (.0174533*REMDIP(I))
        CDIFF = COS (.0174533*(DECELX-REDEC(I)))
        MGTON = REMMAG(I) * 1.0000E0
        DO 50 J = 1, NFPTS
          VANMLY(J,I,2) = 2.0*MGTON*((CDIP*CDIFF*QSUM(J))-(SDIP*PSUM(J)))
          HANMLY(J,I,2) = 2.0*MGTON*((CDIP*CDIFF*PSUM(J))+(SDIP*QSUM(J)))
C*** FOR ANOMALIES SMALL IN COMPARISON TO THE EARTH'S FIELD, THE INTENSITY OF
C*** THE TOTAL ANOMALY DUE TO REMANENT MAGNETIZATION IS THE SUM OF THE
C*** PROJECTIONS OF THE VERTICAL AND HORIZONTAL ANOMALIES DUE TO REMANENT
C*** MAGNETIZATION ALONG THE DIRECTION OF THE EARTH'S FIELD.
          TANMLY(J,I,2) = HANMLY(J,I,2)*CDIPR*CDIFFR + VANMLY(J,I,2)*SDIPR
        50 CONTINUE
C*****END OF ANOMALY CALCULATIONS*****
C*****BEGIN INDIVIDUAL POLYGON OUTPUT SECTION*****
C***PRINT ANOMALIES FOR THIS POLYGON AND PLOT.
      DO 130 I = 1, NPOLY
C ADD REMANENT + INDUCED FOR INDIVIDUAL POLYGONS
        DO 60 J = 1, NFPTS
          VWORK(J) = VANMLY(J,I,1) + VANMLY(J,I,2)
          HWORK(J) = HANMLY(J,I,1) + HANMLY(J,I,2)
          TWORK(J) = TANMLY(J,I,1) + TANMLY(J,I,2)
C*****PRINT ANOMALIES FOR INDIVIDUAL POLYGONS*****
          IF (PRIND) 70,120,70
          70 WRITE (6,6000) I,SUS(I)
6000 FORMAT(1H1,'POLYGON NUMBER ',I2,'/' SUSCEPTIBILITY IS',
          1F10.6,' EMU/' CORNERS ARE ( X , Z )')
          DO 80 J = 1, NSIDES
            80 WRITE (6,7000) X(I,J), Z(I,J)
7000 FORMAT(1H1,'(F8.3,F8.3)')
            M = IFIX (1000.0 * REMMAG(I))
            IF (M) 100,90,100
          90 WRITE (6,8000)
8000 FORMAT (1H0,'MAGNETIZATION IS INDUCED ONLY.')
          GO TO 110
          100 WRITE (6,9000) REMDEC(I),RECDIP(I),REMMAG(I)
9000 FORMAT(1H0,'MAGNETIZATION IS MIXED. REMANENT ',
          1'MAGNETIZATION VECTOR IS DEFINED BY', DEC='F8.3', INC='F8.3', MAGNITUDE='F10.6', EMU')
          110 WRITE (6,9001)
9001 FORMAT(1H0,'ANOMALIES IN GAMMAS',/1X,24('*****'),/,' FIELD POINT'
          1 12X,'INDUCED',29X,'REMANENT',29X,'TOTAL',/17X
          2 3('V',9X,'H',9X,'T',15X),/24('*****'),/
          DO 118 J=1,NFPTS
          118 WRITE (6,9002) FLDPT(J),VANMLY(J,I,1),HANMLY(J,I,1),TANMLY(J,I,1)
          1,VANMLY(J,I,2),HANMLY(J,I,2),TANMLY(J,I,2),VWORK(J),HWORK(J),
          2,TWORK(J)
9002 FORMAT (1H,'F9.3',/3(3F10.3,X))
C***PLOT VERTICAL ANOMALY FOR THIS POLYGON -- REMANENT + INDUCED
      120 IF (PLINDV) 121,123,121
      121 FSCALE = 0.0
        CALL CONARY(VANMLY,I,1,WOR,1,NFPTS)
        CALL CONARY(VANMLY,I,2,WOR,2,NFPTS)
        CALL GRSCAL(NFPTS,WOR,FSCALE,1,XPO*MAX)
        CALL GRSCAL(NFPTS,WOR,FSCALE,1,XPO*MAX)
        CALL GRSCAL(NFPTS,VWORK,FSCALE,1,XPO*MAX)
        WRITE (6,9004) I
9004 FORMAT (1H1,40X,'VERTICAL ANOMALY FOR POLYGON NUMBER ',I2)
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
  CALL SPLOT(NFPTS,FLDPT,WORX1,WORX2,VWORK,FSCALE,REMMAG(I),
    * MAX,IEXP0,ITICKS)
C***PLOT HORIZONTAL ANOMALY FOR THIS POLYGON -- REMANENT + INDUCED
123 IF (PLINH) 124,125,126
124 FSCALE = 0.0
  CALL CONARY(HANMLY,1,1,WOR,1,NFPTS)
  CALL CONARY(HANMLY,1,2,WOR,2,NFPTS)
  CALL GRSCAL(NFPTS,WORX1,FSCALE,IEXP0*MAX)
  CALL GRSCAL(NFPTS,WORX2,FSCALE,IEXP0*MAX)
  CALL GRSCAL(NFPTS,HWORK,FSCALE,IEXP0*MAX)
  WRITE (6,9008) I
9008 FORMAT (1H1,39X,'HORIZONTAL ANOMALY FOR POLYGON NUMBER ',I2)
  CALL SPLOT(NFPTS,FLDPT,WORX1,WORX2,HWORK,FSCALE,REMMAG(I),
    * MAX,IEXP0,ITICKS)
C***PLOT TOTAL ANOMALY FOR THIS POLYGON -- REMANENT + INDUCED
125 IF (PLINT) 126,130,126
126 FSCALE = 0.0
  CALL CONARY(TANMLY,1,1,WOR,1,NFPTS)
  CALL CONARY(TANMLY,1,2,WOR,2,NFPTS)
  CALL GRSCAL(NFPTS,WORX1,FSCALE,IEXP0*MAX)
  CALL GRSCAL(NFPTS,WORX2,FSCALE,IEXP0*MAX)
  CALL GRSCAL(NFPTS,TWORK,FSCALE,IEXP0*MAX)
  WRITE (6,9009) I
9009 FORMAT (1H1,44X,'TOTAL ANOMALY FOR POLYGON NUMBER ',I2)
  CALL SPLOT(NFPTS,FLDPT,WORX1,WORX2,TWORK,FSCALE,REMMAG(I),
    * MAX,IEXP0,ITICKS)
130 CONTINUE
C********END OF INDIVIDUAL POLYGON OUTPUT SECTION*****
C********BEGIN FIELD POINT SUMMATION SECTION*****
C***ZERO ALL FIELD POINT SUMMING ARRAYS.
  DO 135 I = 1,NFPTS
    VIND(I) = 0.0
    VREM(I) = 0.0
    HIND(I) = 0.0
    HREM(I) = 0.0
    TIND(I) = 0.0
    TREM(I) = 0.0
135 CONTINUE
  DO 140 I = 1,NFPTS
    DO 137 J = 1,NPOLY
      VIND(I) = VIND(I) + VANMLY(I,J,1)
      HIND(I) = HIND(I) + HANMLY(I,J,1)
      TIND(I) = TIND(I) + TANMLY(I,J,1)
      VREM(I) = VREM(I) + VANMLY(I,J,2)
      HREM(I) = HREM(I) + HANMLY(I,J,2)
137 TREM(I) = TREM(I) + TANMLY(I,J,2)
      VTOT(I) = VIND(I) + VREM(I)
      HTOT(I) = HIND(I) + HREM(I)
      TTOT(I) = TIND(I) + TREM(I)
140 CONTINUE
C********END FIELD POINT SUMMATION SECTION*****
C********BEGIN COMPOSITE OUTPUT SECTION*****
C***PLOT COMPOSITION FOR VERTICAL ANOMALY DUE TO INDUCTION ONLY
  IF (VCOMP) 145,153,145
145 FSCALE = 0.0
  DO 150 I = 1,NPOLY
    CALL CONARY(VANMLY,1,1,WOR,1,NFPTS)
150 CALL GRSCAL(NFPTS,WORX1,FSCALE,IEXP0*MAX)
    CALL GRSCAL(NFPTS,VIND,FSCALE,IEXP0*MAX)
    WRITE (6,9010) I
9010 FORMAT (1H1,40X,'VERTICAL ANOMALY DUE TO INDUCTION ONLY.')
    CALL CPLOT(NPOLY,NFPTS,FLDPT,VANMLY,1,VIND,FSCALE,1,MAX,IEXP0,
      * ITICKS)
C***PLOT COMPOSITION FOR VERTICAL ANOMALY DUE TO REMANENCE ONLY.
153 IF (VCOMPR) 155,153,155
155 FSCALE = 0.0
  DO 160 I = 1,NPOLY
    CALL CONARY(VANMLY,1,2,WOR,1,NFPTS)
160 CALL GRSCAL(NFPTS,WORX1,FSCALE,IEXP0*MAX)
    CALL GRSCAL(NFPTS,VREM,FSCALE,IEXP0*MAX)
    WRITE (6,9011) I
9011 FORMAT (1H1,40X,'VERTICAL ANOMALY DUE TO REMANENCE ONLY.')
    CALL CPLOT(NPOLY,NFPTS,FLDPT,VANMLY,2,VREM,FSCALE,1,MAX,IEXP0,
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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
* ITICKS)
C***PLOT COMPOSITION FOR HORIZONTAL ANOMALY DUE TO INDUCTION ONLY
163 IF (HCOMP1) 165,173,165
165 FSCALE = 0.0
DO 170 I = 1,NPOLY
CALL CONARY(HANMLY,1,1,WOR,1,NFPTS)
170 CALL GRSCAL(NFPTS,WORK1,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,HIND,FSCALE,IEXPO*MAX)
WRITE (6,9012)
9012 FORMAT (1H1,40X,'HORIZONTAL ANOMALY DUE TO INDUCTION ONLY')
CALL CPLOT(NPOLY,NFPTS,FLDPT,HANMLY,1,HIND,FSCALE*2,MAX,IEXPO,
* ITICKS)
C***PLOT COMPOSITION FOR HORIZONTAL ANOMALY DUE TO REMANENCE ONLY
173 IF (HCOMP2) 175,183,175
175 FSCALE = 0.0
DO 180 I = 1,NPOLY
CALL CONARY(HANMLY,1,2,WOR,1,NFPTS)
180 CALL GRSCAL(NFPTS,WORK1,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,HREM,FSCALE,IEXPO*MAX)
WRITE (6,9013)
9013 FORMAT (1H1,40X,'HORIZONTAL ANOMALY DUE TO REMANENCE ONLY')
CALL CPLOT(NPOLY,NFPTS,FLDPT,HANMLY,2,HREM,FSCALE*2,MAX,IEXPO,
* ITICKS)
C***PLOT COMPOSITION FOR TOTAL ANOMALY DUE TO INDUCTION ONLY
183 IF (TCOMP1) 185,193,185
185 FSCALE = 0.0
DO 190 I = 1,NPOLY
CALL CONARY(TANMLY,1,1,WOR,1,NFPTS)
190 CALL GRSCAL(NFPTS,WORK1,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,TIND,FSCALE,IEXPO*MAX)
WRITE (6,9014)
9014 FORMAT (1H1,40X,'TOTAL ANOMALY DUE TO INDUCTION ONLY')
CALL CPLOT(NPOLY,NFPTS,FLDPT,TANMLY,1,TIND,FSCALE*3,MAX,IEXPO,
* ITICKS)
C***PLOT COMPOSITION FOR TOTAL ANOMALY DUE TO REMANENCE ONLY
193 IF (TCOMP2) 195,205,195
195 FSCALE = 0.0
DO 200 I = 1,NPOLY
CALL CONARY(TANMLY,1,2,WOR,1,NFPTS)
200 CALL GRSCAL(NFPTS,WORK1,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,TREM,FSCALE,IEXPO*MAX)
WRITE (6,9015)
9015 FORMAT (1H1,40X,'TOTAL ANOMALY DUE TO REMANENCE ONLY')
CALL CPLOT(NPOLY,NFPTS,FLDPT,TANMLY,2,TREM,FSCALE*3,MAX,IEXPO,
* ITICKS)
205 CONTINUE
C*****END COMPOSITE OUTPUT SECTION*****
C*****BEGIN ANOMALIES DUE TO ALL POLYGONS OUTPUT SECTION*****
C***PRINT OUT ANOMALIES DUE TO ALL POLYGONS
IF (PRINT1) 206,260,206
206 REMFLG = 0.0
DO 210 I = 1,NPOLY
210 REMFLG = REMFLG + REMMAG(I)
WRITE (6,9016) NPOLY
9016 FORMAT (1H1,'ANOMALIES DUE TO ALL ',I2,' POLYGON(S)')
IF (REMFLG) 230,220,230
220 WRITE (6,8000)
GO TO 240
230 WRITE (6,9017)
9017 FORMAT (1H0,'MAGNETIZATION IS MIXED.')
240 WRITE (6,9001)
DO 250 I = 1,NFPTS
250 WRITE (6,9002) FLDPT(I),VIND(I),HIND(I),TIND(I),VREM(I),HREM(I),
1 TREM(I),VTOT(I),HTOT(I),TTOT(I)
C***PLOT VERTICAL ANOMALY DUE TO ALL POLYGONS -- REMANENT + INDUCED
260 IF (PLVAL) 270,281,270
270 FSCALE = 0.0
CALL GRSCAL(NFPTS,VIND,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,VREM,FSCALE,IEXPO*MAX)
CALL GRSCAL(NFPTS,VTOT,FSCALE,IEXPO*MAX)
WRITE (6,9018)
9018 FORMAT (1H1,40X,'VERTICAL ANOMALY DUE TO ALL POLYGONS')
CALL SPLOT(NFPTS,FLDPT,VIND,VREM,VTOT,FSCALE,REMFLG*MAX,IEXPO,
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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  * ITICKS)
C*****PLOT HORIZONTAL ANOMALY DUE TO ALL POLYGONS - REMANENT + INDUCED
280 IF (PLHAL) 290,300,290
290 FSCALE = 0.0
    CALL GRSCAL(NFPTS,HIND,FSCALE,IEXPO,MAX)
    CALL GRSCAL(NFPTS,HREM,FSCALE,IEXPO,MAX)
    CALL GRSCAL(NFPTS,HTOT,FSCALE,IEXPO,MAX)
    WRITE (6,9019)
9019 FORMAT (1H1,41X,'HORIZONTAL ANOMALY DUE TO ALL POLYGONS')
    CALL SPLOT (NFPTS,FLDPT,HIND,HRE,HTOT,FSCALE,REMFLG,MAX,IEXPO,
  * ITICKS)
C*****PLOT TOTAL ANOMALY DUE TO ALL POLYGONS- REMANENT + INDUCED
300 IF (PLTAL) 310,315,310
310 FSCALE = 0.0
    CALL GRSCAL(NFPTS,TIND,FSCALE,IEXPO,MAX)
    CALL GRSCAL(NFPTS,TREM,FSCALE,IEXPO,MAX)
    CALL GRSCAL(NFPTS,TTOT,FSCALE,IEXPO,MAX)
    WRITE (6,9020)
9020 FORMAT (1H1,45X,'TOTAL ANOMALY DUE TO ALL POLYGONS')
    CALL SPLOT (NFPTS,FLDPT,TIND,TREM,TTOT,FSCALE,REMFLG,MAX,IEXPO,
  * ITICKS)
315 GO TO 1
320 CALL EXIT
    END
1234567890123456789012345678901234567890123456789012345678901234567890

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1234567890123456789012345678901234567890123456789012345678901234567890
  SUBROUTINE GRSCAL(NFPTS,A,FSCALE,IEXPO,MAX)
    DIMENSION A(200)
C*****
C THIS SUBROUTINE FINDS THE VARIABLE WITH THE GREATEST ABSOLUTE VALUE IN THE
C ARRAY A WHICH HAS NFPTS ELEMENTS.
C FSCALE IS THEN SET EQUAL TO THE SMALLEST VALUE FOR FULL SCALE WHICH WILL
C BEST PRESENT THE ELEMENTS OF A.
C IEXPO AND MAX ARE SUCH THAT FSCALE = MAX * 10 ** IEXPO.
C*****
    DO 15 I = 1,NFPTS
      5 IF (ABS(A(I))-FSCALE) 15,15,10
    10 FSCALE = ABS(A(I))
    15 CONTINUE
C** IF ALL VALUES OF THE ARRAY ARE ZERO, SET FSCALE = 1.0 ARBITRARIY,
    IF (FSCALE) 20,20,25
    20 FSCALE = 1.0
    25 EXPO = LOG10 (FSCALE)
    IF (EXPO) 30,35,40
    30 IEXPO = EXPO + 1
    GO TO 45
    35 MAX = 1
    RETURN
    40 IEXPO = EXPO
    45 J = FSCALE/(10.0**IEXPO)
    IF (J.LT.2) GO TO 50
    IF (J.LT.5) GO TO 55
    FSCALE = 10.0 * (10.0**IEXPO)
    MAX = 10
    RETURN
    50 FSCALE = 2.0 * (10.0**IEXPO)
    MAX = 2
    RETURN
    55 FSCALE = 5.0 * (10.0**IEXPO)
    MAX = 5
    RETURN
    END
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
SUBROUTINE POLYPO (NFPTS, NSIDES, FLDPT, OHSVHT, PSUM, QSUM,
* INDEX, X, Z, DELX)
  DIMENSION FAX(21), ZEE(21), FLDPT(200), PSUM(200), QSUM(200)
  DIMENSION X(10,20), Z(10,20)
  C***,*****
  C***READ COORDINATE CARDS FOR POLYGON CORNERS
  C***ONE COORDINATE CARD FOR EACH CORNER OF THE POLYGON--CLOCKWISE ORDER.
  C***FXX(I) = X COORDINATE OF THE ITH CORNER OF POLYGON
  C***ZEE(I) = Z COORDINATE OF THE ITH CORNER OF POLYGON, POSITIVE DOWN.
  C***,*****
  DO 10 I = 1, NSIDES
    READ (5,1000) EXX(I), ZEE(I)
  1000 FORMAT( )
  C***SAVE CORNER COORDINATES FOR PRINTOUT.
  X(INDEX,1) = EXX(1)
  Z(INDEX,1) = ZEE(1)
  10 NSIDP1 = NSIDES + 1
  EXX(NSIDP1) = EXX(1)
  ZEE(NSIDP1) = ZEE(1)
  DO 150 I = 1, NFPTS
    PSUM(I) = 0.0
    QSUM(I) = 0.0
    X1 = EXX(1) - FLDPT(1)
    Z1 = ZEE(1) + OHSVHT
    15 RSQ1 = (X1*X1 + Z1*Z1)
  C***IF X AND Z ARE BOTH ZERO, ATAN2 GIVES ERROR, SO CHANGE Z SLIGHTLY
  IF (RSQ1.NE.0.0) GO TO 17
  Z1 = .0001 * DELX
  THETD = 0.0
  GL = ALOG (.0001 * DELX * DELX)
  GO TO 15
  17 THETA = ATAN2(Z1,X1)
  J = 2
  20 X2 = EXX(J) - FLDPT(1)
  Z2 = ZEE(J) + OHSVHT
  25 RSQ2 = (X2*X2 + Z2*Z2)
  IF (RSQ2.NE.0.0) GO TO 27
  Z2 = .0001 * DELX
  THETD = 0.0
  GL = ALOG (.0001 * DELX * DELX)
  GO TO 25
  27 THETB = ATAN2(Z2,X2)
  IF (Z1-Z2) 40,30,40
  30 P = 0.0
  Q = 0.0
  GO TO 120
  40 OMEGA = THETA - THETB
  IF (OMEGA) 60,50,50
  50 IF (OMEGA-3.1415927) 70,70,60
  60 IF (OMEGA+3.1415927) 80,70,70
  70 THETD = OMEGA
  GO TO 110
  80 IF (OMEGA) 90,100,100
  90 THETD = OMEGA - 6.2831853
  GO TO 110
  100 THETD = OMEGA + 6.2831853
  110 GL = 0.5 * ALOG(RSQ2/RSQ1)
  115 X12 = X1 - X2
  Z21 = Z2 - Z1
  XSQ = X12 * X12
  ZSQ = Z21 * Z21
  XZ = Z21 * X12
  P = -((ZSQ/(XSQ+ZSQ))*THETD) + ((XZ/(XSQ+ZSQ))*GL)
  Q = -(THETD*(XZ/(XSQ+ZSQ))) - (GL*(ZSQ/(XSQ+ZSQ)))
  120 PSUM(I) = PSUM(I) + P
  QSUM(I) = QSUM(I) + Q
  C***RELABEL VARIABLES INVOLVING ONLY THE SECOND POLYGON CORNER AS THE
  C*** VARIABLES INVOLVING ONLY THE FIRST POLYGON CORNER SO THEY DON'T
  C*** HAVE TO BE CALCULATED AGAIN.
  140 X1 = X2
  Z1 = Z2
  RSQ1 = RSQ2
  THETA = THETB
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
C***CHECK TO SEE IF ALL SIDES HAVE BEEN DONE.
      J = J + 1
      JR = J - 1
      IF(JR-NSIDP1) 20,150,150
150 CONTINUE
      RETURN
      END
1234567890123456789012345678901234567890123456789012345678901234567890

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1234567890123456789012345678901234567890123456789012345678901234567890
      SUBROUTINE CONARY (ARRAY1,I,J,ARRAY2,NFPTS)
      DIMENSION ARRAY1(200,10,2),ARRAY2(200)
C***,*****
C   THIS SUBROUTINE SELECTS THE DESIGNATED ELEMENTS OF THE THREE DIMENSIONAL
C   ARRAY, ARRAY1 AND PUTS THEM INTO A ONE DIMENSIONAL ARRAY, ARRAY2)
C   THAT IS ARRAY2(I) = ARRAY1(I,I,J) THRU TO
C   ARRAY2(NFPTS) = ARRAY1(NFPTS,I,J) ARE PERFORMED.
C***,*****
      DO 10 K = 1,NFPTS
      10 ARRAY2(K) = ARRAY1(K,I,J)
      RETURN
      END
1234567890123456789012345678901234567890123456789012345678901234567890

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Table 6. (Continued)

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1234567890123456789012345678901234567890123456789012345678901234567890
SUBROUTINE SPLOT (NFPTS,FLOPT,A,B,C,FSCALE,REMFLG,MAX,IEXPO,
* ITICKS)
  DIMENSION A(200),B(200),C(200),FLOPT(200),ALINE(101)
  REAL I1
  DATA BLNK/1H /,AST/1H*,DOT/1H./,RR/1HR/,II/1H/,DASH/1H-/ ,PLS/1H+
C*****
C THIS SUBROUTINE GENERATES A LINE PRINTER PLOT OF THE ELEMENTS OF THE ARRAYS
C A , B , AND C VS. THE FIELD POINTS, FLOPT .
C FLOPT ELEMENTS ARE LISTED ALONG THE ABSCISSA AND THE ARRAY ELEMENTS
C PLOTTED ALONG THE ORDINATE. HATCH MARKS ARE PLACED ALONG THE ABSCISSA
C EVERY ITICKS TH POINT.
C FSCALE IS THE FULL SCALE LIMIT OF THE PLOT.
C FSCALE = MAX * 10 ** IEXPO
C REMFLG PERMITS CHOICE OF PLOTTING SYMBOLS DEPENDING ON WHETHER THE MAG-
C NETIZATION IS MIXED OR NOT.
C REMFLG 0 , USE SYMBOLS FOR INDUCTION ONLY
C      = 1 , USE SYMBOLS FOR MIXED MAGNETIZATION
C*****
  N = ITICKS
  WRITE (6,500) IEXPO,IEXPO,MAX,MAX
500  FORMAT (1H ,17X,I1,99X,I1,/,12X,'-',I2,1X10',47X,'0',46X,'+',
1  I2,1X10',/, FIELD POINT',3X,21(' ',),/15X,21(5H' '))
  WRITE (6,600)
600  FORMAT (1H+,14X,20('+. ...'),'+')
  DO 10 I = 1,101
    ALINE(I) = BLNK
    ALINE( 1) = DOT
    ALINE( 51) = DOT
    ALINE(101) = DOT
    DO 30 J = 1,NFPTS
      IF (ITICKS-N) 13,12,13
12  N = 1
      ALINE( 1) = PLS
      ALINE( 2) = DASH
      ALINE( 50) = DASH
      ALINE( 51) = PLS
      ALINE( 52) = DASH
      ALINE(100) = DASH
      ALINE(101) = PLS
      GO TO 14
13  N = N + 1
14  J = 51,500001 + (A(I)/FSCALE)*50
      M = IFIX (1000.0 * REMFLG)
      IF (M) 16,15,16
15  ALINE(J) = AST
      L = J
      K = J
      GO TO 18
16  ALINE(J) = II
      K = 51,500001 + (B(I)/FSCALE) * 50.0
      ALINE(K) = RR
      L = 51,500001 + (C(I)/FSCALE) * 50.0
      ALINE(L) = AST
18  WRITE (6,1000) FLOPT(I),ALINE
1000 FORMAT(1H ,F10.2,4X,101A1)
      WRITE (6,2000)
2000 FORMAT (1H+,14X,1H,99X,1H,+)
      IF (N-1) 25,20,25
20  ALINE( 2) = BLNK
      ALINE( 50) = BLNK
      ALINE( 52) = BLNK
      ALINE(100) = BLNK
25  ALINE(L) = BLNK
      ALINE(K) = BLNK
      ALINE(J) = BLNK
      ALINE( 1) = DOT
      ALINE( 51) = DOT
      ALINE(101) = DOT
30  CONTINUE
      WRITE (6,3000)
3000 FORMAT (1H ,14X,20('+. ...'),'+')
      RETURN
  END
1234567890123456789012345678901234567890123456789012345678901234567890

```

Table 6. (Concluded)

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1234567890123456789012345678901234567890123456789012345678901234567890
SUBROUTINE CPLOT(NPOLY,NFPTS,FLDPT,A,K,B,FSCALE,INDEX,MAX,IEXP0,
* ITICKS)
  DIMENSION FLDPT(200),A(200,10,2),B(200),ALINE(101),SUM(3),LINE(11)
  REAL NUM(10)
  DATA BLNK/1H /,DOT/1H./,(NUM(I),I=1,10)/1H1,1H2,1H3,1H4,1H5,1H6,
* 1H7,1H8,1H9,1H0/,(SU(I),I=1,3)/1HV,1HH,1HT/,DASH/1H-/ ,PLS/1H+/
C THIS SUBROUTINE GENERATES A COMPOSITE LINE PRINTER PLOT OF THE ELEMENTS OF
C THE ARRAY A AS FOLLOWS. NOTE A IS DIMENSIONED AS A(200,10,2).
C THIS SUBROUTINE WILL PLOT THE POINTS A(ALPHA,BETA,1) OR
C A(ALPHA,BETA,2) AS DESCRIBED BY THE FOLLOWING. NPOLY PLOTTING SYMBOLS
C ARE PLOTTED FOR EACH VALUE OF THE ABSCISSA, FOR NFPTS FIELD POINTS. ALSO
C FOR EACH VALUE OF THE ABSCISSA A PLOTTING SYMBOL REPRESENTING THE CORRESPOND-
C ING ELEMENT OF THE ARRAY B, WHICH IN THIS CASE REPRESENTS THE ALGEBRAIC
C SUM OF THE NPOLY VALUES FOR ARRAY A IS ALSO PLOTTED.
C FLDPNT ELEMENTS ARE LISTED ALONG THE ABSCISSA AND THE ARRAY ELEMENTS OF
C A AND B ARE PLOTTED ALONG THE ORDINATE. MATCH MARKS ARE PLACED ALONG
C THE ABSCISSA EVERY ITICKS TH POINT. FSCALE IS THE FULL SCALE LIMIT OF
C THE PLOT. FSCALE = MAX * 10 ** IEXP0.
  N = ITICKS
  WRITE (6,500) IEXP0,IEXP0,MAX,MAX
500  FORMAT (1H,17X,11.99X,11,/,12X,11.99X,12,1X10',47X,10',46X,11,/,
1 12,1X10',/,1 FIELD POINT',3X,21(' ',/),15X,21(5H' '))
  WRITE (6,600)
600  FORMAT (1H,14X,20('+. ....'),++)
  DO 10 I = 1,101
  10  ALINE(I) = BLNK
  ALINE( 1) = DOT
  ALINE( 51) = DOT
  ALINE(101) = DOT
  DO 40 I = 1,NFPTS
  IF (ITICKS-N) 13,12,13
  12  N = 1
  ALINE( 1) = PLS
  ALINE( 2) = DASH
  ALINE( 50) = DASH
  ALINE( 51) = PLS
  ALINE( 52) = DASH
  ALINE(100) = DASH
  ALINE(101) = PLS
  GO TO 14
  13  N = N + 1
  14  DO 20 J = 1,NPOLY
  L = 51.500001 + (A(I,J,K)/FSCALE) * 50.0
C***SAVE NUMBER OF THE ELEMENT WHICH WAS CHANGED FROM BLANK
  LINE(J) = L
C***REPLACE BLANK LINE ELEMENT BY CHARACTER FOR DATA POINT.
  20  ALINE(L) = NUM(J)
  L = 51.500001 + (B(I)/FSCALE) * 50.0
  ALINE(L) = SUM(INDEX)
C***J IS NOW = NPOLY + 1
  LINE(J) = L
  WRITE(6,1000) FLDPT(I),ALINE
1000  FORMAT(1H,1F10.2,4X,101A1)
  WRITE (6,2000)
2000  FORMAT (1H,14X,1H,99X,1H,/)
C***RESTORE BLNKS TO THOSE LINE ELEMENTS WHICH WERE CHANGED
  DO 30 L = 1,J
  M = LINE(L)
  30  ALINE(M) = BLNK
C***RESTORE DOTS TO THOSE LINE ELEMENTS WHICH MAKE UP THE AXIS AND MAY HAVE
C*** BEEN CHANGED.
  IF (N-1) 38,35,38
  35  ALINE( 2) = BLNK
  ALINE( 50) = BLNK
  ALINE( 52) = BLNK
  ALINE(100) = BLNK
  38  ALINE( 1) = DOT
  ALINE( 51) = DOT
  ALINE(101) = DOT
  40  CONTINUE
  WRITE (6,3000)
3000  FORMAT (1H,14X,20('+. ....'),++)
  RETURN
  END
1234567890123456789012345678901234567890123456789012345678901234567890

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## APPENDIX IV

## PALEOMAGNETIC DATA COLLECTION AND REDUCTION

Samples from the Meriwether dike were collected and analyzed for NRM (Natural Remanent Magnetization) by Doyle Watts of the Ohio State University. Thirteen cores were obtained from a single outcrop of the Meriwether dike 5.5 miles northeast of Greenville on Georgia State Highway 362. In all, 26 one inch cylinders were cut from the cores. The samples were analyzed for Natural Remanent Moment (no magnetic cleaning) using a Schonstedt SSML Spinner Magnetometer. Direction and magnitude of the NRM are given for each sample in Table 7. Sample numbers are those used by Watts (personal communication) and the letter A, B, or C following the number indicates the first, second, or third cylinder cut from a core. Cores NW73174 through NW73182 were taken from the center portion of the dike and the remaining cores from the chilled edges of the dike.

Table 7. Natural Remanent Moments of Cores Taken From  
The Meriwether Dike (Watts, Personal Communication)

Ohio State University Sample Number	Remanent Moment		
	Magnitude	Declination	Inclination
NW73174B	0.00144	7.82	26.26
NW73175A	0.00144	2.98	28.54
NW73175B	0.00156	8.86	35.49
NW73176B	0.00150	1.03	35.39
NW73177B	0.00150	21.69	27.86
NW73177C	0.00151	18.31	35.96
NW73178B	0.00174	12.42	25.20
NW73178C	0.00168	13.94	26.29
NW73179B	0.00152	18.10	22.61
NW73179C	0.00149	22.29	21.84
NW73180B	0.00173	10.08	26.39
NW73180C	0.00175	14.89	25.46
NW73180D	0.00170	12.01	27.20
NW73181B	0.00166	14.77	29.96
NW73182B	0.00162	16.20	23.21
NW73182C	0.00153	19.51	25.58
NW73183B	0.00215	36.16	27.72
NW73183C	0.00220	36.84	31.15
NW73184B	0.00211	33.78	23.38
NW73184C	0.00173	30.42	16.23
NW73185B	0.00201	42.87	27.96
NW73185C	0.00193	46.15	27.09
NW73186B	0.00220	42.11	32.70
NW73186C	0.00210	46.66	31.44
NW73187B	0.00212	49.75	30.89
NW73187C	0.00216	42.37	30.95

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